GEOLOGIC SETTING AND WATER QUALITY OF SELECTED BASINS
IN THE ACTIVE COAL-MINING AREAS OF OHIO,
JUNE 1985 THROUGH DECEMBER 1986

By Allison L. Jones

U.S. GEOLOGICAL SURVEY

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#### CONVERSION FACTORS

For use of readers who prefer to use metric (International System) units rather than the inch-pound terms used in this report, the following conversion factors may be used:

Multiply inch pound unit	<u>By</u>	To obtain metric unit
foot (ft)	0.3048	meter (m)
foot per mile (ft/mi)	0.1894	<pre>meter per kilometer   (m/km)</pre>
cubic foot per second (ft <sup>3</sup> /s)	0.02832	cubic meter per second (m <sup>3</sup> /s)
<pre>gallon per minute   (gal/min)</pre>	0.06309	liter per second (L/s)
inch (in.) square mile (mi <sup>2</sup> )	25.4 2.590	millimeter (mm) square kilometer (km²)
<pre>foot per mile (ft/mi)  cubic foot per second   (ft<sup>3</sup>/s)  gallon per minute   (gal/min) inch (in.)</pre>	0.1894 0.02832 0.06309 25.4	meter per kilometer (m/km) cubic meter per seco (m <sup>3</sup> /s) liter per second (L/s) millimeter (mm)

Temperature in degrees Fahrenheit  $({}^{\circ}F)$  can be converted to degrees Celsius  $({}^{\circ}C)$  as follows:

$$^{\circ}F = 1.8 \times ^{\circ}C + 32$$

<u>Sea level</u>: In this report "sea level" refers to the National Geodetic Vertical Datum of 1929 (NGVD of 1929)—a geodetic datum derived from a general adjustment of the first-order level nets of both the United States and Canada, formerly called "Mean Sea Level of 1929."

# GEOLOGIC SETTING AND WATER QUALITY OF SELECTED BASINS IN THE ACTIVE COAL-MINING AREAS OF OHIO, JUNE 1985 THROUGH DECEMber 1986

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#### ABSTRACT

During the initial 18 months of a 7-year study, water-quality data were collected from 20 basins in the active coalmining areas of Ohio. The study area is mostly within the unglaciated eastern part of Ohio along the western edge of the Appalachian Plateaus physiographic province. The area is underlain by sandstone, shale, coal, and limestone of Mississippian, Pennsylvanian, and Permian age.

One to three long-term surface-water sites were assigned to each basin to document present water quality and changes in water quality over time. A total of 40 sites were sampled twice yearly at low flow. This report contains the results of the first three rounds of sampling. In addition, two to three individual basins are being chosen for intensive study during each year of the project. Additional data from the first five basins are included in this report.

For each of these intensively studied basins, 10 additional surface-water sites were selected for sampling to represent the water quality of the basin in more detail than was possible with one to three samples. Where present, a productive aquifer (that is, an aquifer capable of producing enough water for commercial, light industrial, or public water supplies) was investigated by sampling three to five wells and measuring water levels concurrently. The purpose of this ground-water sampling is to describe any productive aquifer present in areas where coal-mining permits may be issued.

Specific-conductance values for 100 samples ranged from 270 to 2,800  $\mu$ S/cm. One hundred samples from the long-term surfacewater network had a median specific conductance of 760  $\mu$ S/cm (microsiemens per centimeter at 25 °C) and a range of 270 to 2,800  $\mu$ S/cm (microsiemens per centimeter at 25 °C) and had a median of 760  $\mu$ S/cm. pH values ranged from 2.70 to 8.85 and had a median of 7.70. Alkalinity values ranged from 0 to 390 mg/L (milligrams per liter) as CaCO<sub>3</sub> and had a median of 122 mg/L.

The initial five basins chosen for intensive study were the Stillwater Creek basin; Symmes, Ice, and Indian Guyan Creeks basin; Moxahala Creek basin; Little Beaver Creek basin; and McMahon, Captina, and Sunfish Creeks basin. In the Stillwater Creek basin, specific conductance ranged from 390 to 3,300 µS/cm, pH ranged from 7.00 to 8.00, and alkalinity ranged from 78 to 372 mg/L as CaCO3. In the Symmes, Ice, and Indian Guyan Creeks basin, the specific conductance ranged from 250 to 465 µS/cm, the pH ranged from 6.70 to 7.90, and the alkalinity ranged from 10 to 182 mg/L as CaCO3. In the Moxahala Creek basin, specific conductance ranged from 520 to 4,600 µS/cm, pH ranged from 3.44 to 9.20, acidity ranged from 9.0 to 465 mg/L as CaCO3, and alkalinity ranged from 2 to 425 mg/L as CaCO3. In the Little Beaver Creek basin, specific conductance ranged from 330 to 1,210 µS/cm, pH ranged from 7.49 to 8.83, and alkalinity ranged from 35 to 257 mg/L as CaCO3. In the McMahon, Captina, and Sunfish Creeks basin, specific conductance ranged from 350 to 1,200 µS/cm, pH ranged from 7.94 to 8.78, and alkalinity ranged from 85 to 192 mg/L as CaCO<sub>3</sub>.

Chemical analyses for 150 surface-water samples and 17 ground-water samples are presented. At the completion of the 7-year study a large amount of baseline data will have been gathered for use by the Ohio Department of Natural Resources to facilitate permit application and compliance and to study water-quality trends.

#### INTRODUCTION

#### Background

Coal has been mined in the eastern third of Ohio since 1804. The first mines were small pick-and-shovel operations in which coal was removed by hand (Eavenson, 1942). Drift mines, tunneled into the coal seams, were developed next. As demand for coal steadily increased, seams too deep for drift mining were reached by vertical shafts (Pfaff and others, 1981). Surface mining began on a small scale in Ohio in 1913. As the supply of easily obtainable coal in the deep coal mines dwindled and its removal became less efficient and less economical, there was a shift to surface extraction. In 1986, 59 percent of all coal produced in Ohio was surface mined (Ohio Department of Natural Resources, 1987).

Coal mining has affected the environment for many years. Only recently have stringent mining and reclamation laws been enacted to help restore and protect the land and water. A need remains to better understand the cumulative effects of coal mining on the quality of water resources in Ohio. Surface-water-quality data have been collected in the past, although over a relatively short period, but data collection has concentrated in areas of abandoned mines. Ground-water-quality data are scarce.

The U.S. Geological Survey began a 7-year study in 1985 in cooperation with the Ohio Department of Natural Resources (ODNR), Division of Reclamation, to obtain baseline hydrologic data for ODNR's evaluation of surface-mining permit application and compliance. The permitting process is designed to prevent future mining from adversely affecting water supplies on which communities and industries depend.

## Purpose and Scope

This report presents the results of the first 18 months of the 7-year study (June 1985 through December 1986). The purpose of this report is to (1) describe the hydrogeologic and physiographic settings of the study areas; (2) present surface-water-quality data for 20 actively mined basins in eastern Ohio; and (3) describe the ground-water quality of a productive aquifer, where present, in each of the five basins selected for intensive study. Trends will not be investigated until further data are collected.

Data were obtained from a network of 40 long-term surface—water sites throughout the 20 basins (fig. 1, in pocket), each of which was sampled twice yearly at low flow. In the five intensively studied basins, an additional 10 surface—water sites (short-term) were sampled during the fall low-flow season.

## Description of the Study Area

The study area includes all or part of 29 counties located in the coal-bearing area of eastern Ohio (fig. 2, in pocket). This area is mostly within the unglaciated part of Ohio along the western edge of the Appalachian Plateaus physiographic province (Fenneman, 1938), where local relief is as much as 500 feet (Razem and Sedam, 1985). A map characterizing land use of eastern Ohio is shown in figure 3. The entire coal-producing area (approximately 8,350 square miles) is drained by the Ohio River and contributes an average annual runoff of 14 inches (Ohio Department of Natural Resources, 1962).

The study area has been divided into 20 basins, which are listed in downstream order in table 1 and shown in figure 2. The five basins selected for intensive study were the Stillwater Creek basin; Symmes, Ice, and Indian Guyan Creeks basin; Little Beaver Creek basin; Moxahala Creek basin; and McMahon, Captina, and Sunfish Creeks basin.

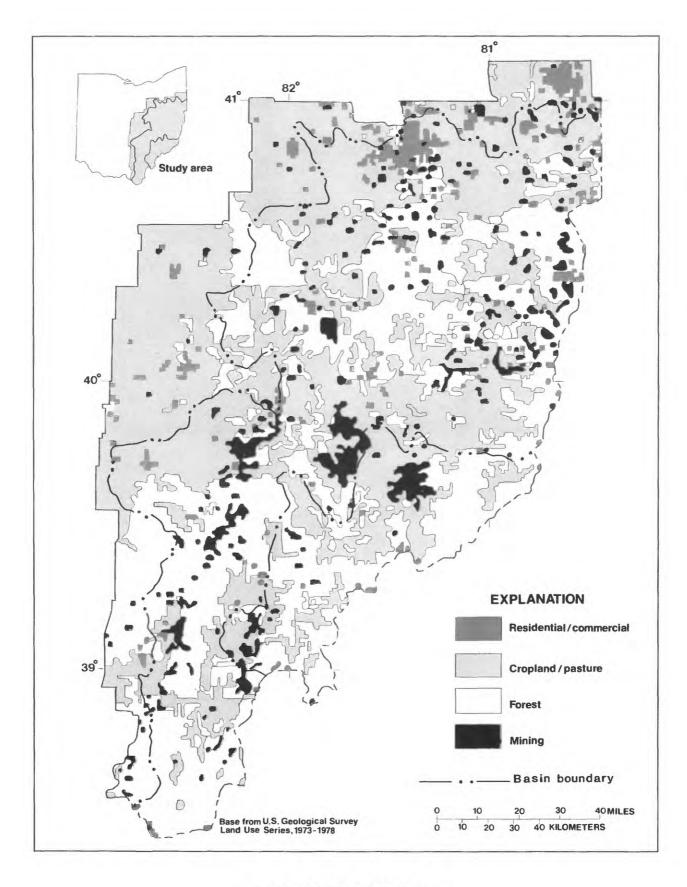


Figure 3.--Land use of eastern Ohio.

# Table 1.--Study basins, in downstream order, and assigned identification code

Basin I.D.	Basin name							
A	Little Beaver Creek							
В	Yellow/Cross Creeks							
С	Short/Wheeling Creeks							
D	McMahon/Captina/Sunfish Creeks							
E	Sandy Creek							
F	Conotton Creek							
G	Middle Tuscarawas River/Sugar Creek							
H	Stillwater Creek							
I	Lower Tuscarawas River							
J	Walhonding River							
K	Upper Wills Creek							
L	Lower Wills Creek							
М	Upper Muskingum River							
N	Middle Muskingum River							
0	Moxahala Creek							
P	Upper Hocking River							
Q	Middle Hocking River							
R	Upper Raccoon Creek							
S	Lower Raccoon Creek							
T	Symmes/Ice/Indian Guyan Creek							

The climate is characterized by moderate extremes of humidity and temperature. Mean daily temperatures range from 19.9 °F in January to 80.0 °F in July to the north, and from 28.0 °F to 90.0 °F to the south (Pierce, 1959). Precipitation is greatest in early spring and least in the fall and averages 40 inches per year. Figure 4 shows a contour map of the average annual precipitation in Ohio (Ohio Department of Natural Resources, 1962).

Coal beds are found in a 32-county area of eastern Ohio (Brant and Delong, 1960). The area is underlain by rocks of Mississippian, Pennsylvanian, and Permian age. Rock types are usually present in sequences of (with increasing depth) coal, limestone, calcareous shale, sandstone, and limestone (Brant and Moulton, 1960).

The outcrop pattern (fig. 5) from west to east progresses from older to younger units, which trend north-northeast and dip regionally to the southeast at approximately 30 feet per mile toward the Appalachian basin. The regional trend of the Pennsylvanian System is modified locally by numerous low structural features (Lamborn, 1951).

The oldest formation of Pennsylvanian age is the Pottsville Formation (fig. 6), in which conglomeratic sandstones are dominant. Above this are the Allegheny, Conemaugh, and Monongahela Formations, all of Pennsylvanian age, which comprise alternating beds of shale, sandstone, coal, and thin limestones. The Dunkard Group of Pennsylvanian and basal Permian age is a variable series of rocks composed of beds of red shale (which is the most abundant rock type), limestone, sandstone, and coal. In the south, the limestone and coal are scarce, and the series consists chiefly of shale and sandstone (Collins, 1979).

The proportion of sandstone strata increases with age. The Allegheny Formation is 40 percent sandstone, and the remainder is composed of shale and clay. The Conemaugh Formation is not more than 30 percent sandstone, and the Monongahela Formation consists of shale, limestone, and not more than 15 percent sandstone (Stout and others, 1943). The carbonate content decreases with age in the Pennsylvanian rocks; therefore, the Monongahela Formation contains more carbonate rocks than the Allegheny and Conemaugh Formations (Razem and Sedam, 1985). Deposits of unconsolidated silts, nearly 100 feet thick in some places, are found chiefly in the bottoms of old valleys that now have no major drainage, and in the lower parts of many smaller valleys. In the main valleys, these silts have largely been removed, and the valleys now are filled with sands and gravels, which have been terraced by the present streams. In the valleys that have carried glacial waters (but have not been glaciated), these coarse, fluvial deposits are common and are composed chiefly of glacial outwash (Happ, 1934).

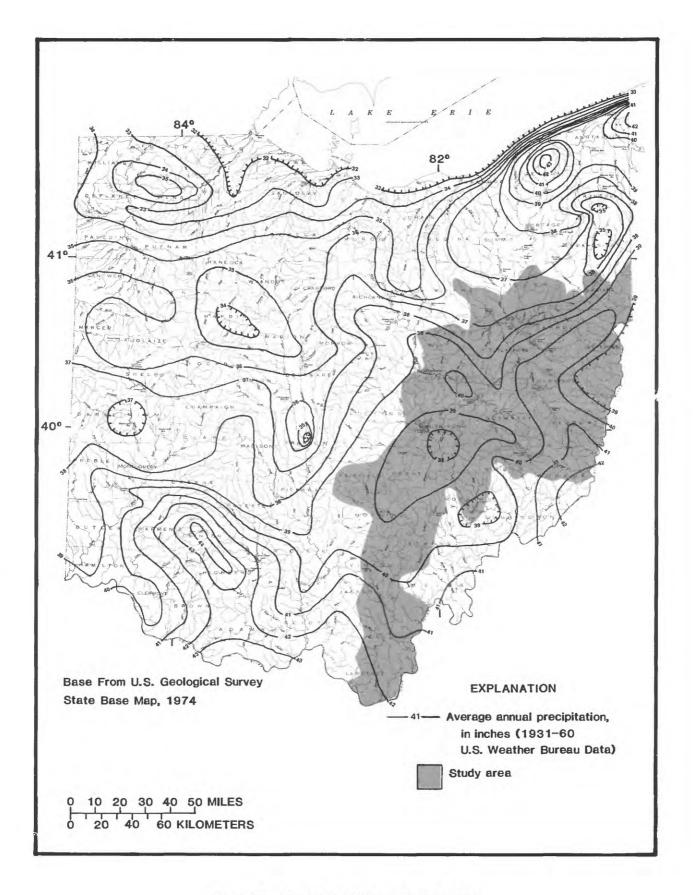


Figure 4.--Average annual precipitation in Ohio.

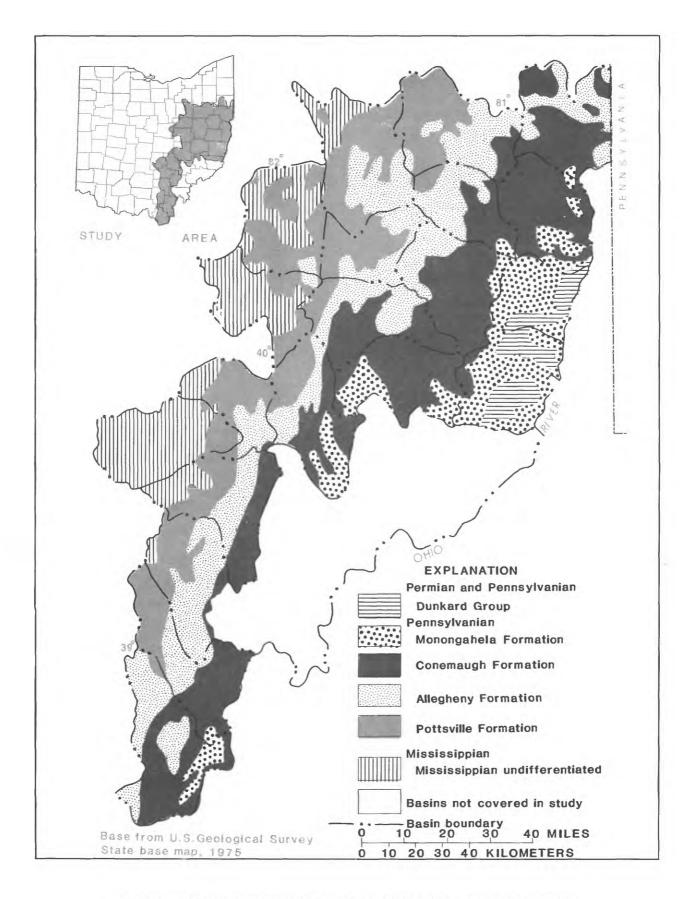


Figure 5.--Surficial geology of study area. (Modified from Collins, 1979).

System	Gro For	ma-	Description	Important coal beds				
Permian	Group	Greene	Mostly red shales and thin limestones, localized coals and sandstone bodies. Present only in small areas.	No.	12	Washington		
Pennsylva- nian and Permian	Dunkard (	Washington	Sandstones, shales, and minor coals. Sandstones are typically micaceous, fine to medium grained, and have thin conglomeratic zones. Locally, sandstones may be massively developed.					
	shale, sandstone, and limestone. Sandstones tend to be fine to medium grained, micaceous, and patchy in devel ment. Compared to other Pennsylvanian units, the Monongahela has a smaller proportion of sandstone and larger proportion of limestone. Limestones tend to be		Monongahela has a smaller proportion of sandstone and a larger proportion of limestone. Limestones tend to be marly, freshwater types. Secondary porosity along frac-	No. No. No.	10 9			
	4	Thick repetitious succession of shales and patchy sand- stones interspersed with thin, discontinuous coals and clays and widespread limestones. The lower limestones are of marine origin, whereas those in the upper part are marly, freshwater types. Secondary porosity along fractured surfaces is well developed locally.						
Pennsylvanian	4	Villegillelik	Repetitious succession of important coal-bearing strata interspersed with several fine to coarse-grained massive, cross-bedded sandstones and thin, persistent limestones. Sandstones, though widespread, have considerable local lateral variation. Solution cavities are developed locally in the limestones.	No. No. No. No.	6 5 4A	Upper Freeport Lower Freeport Middle Kittannin Lower Kittanning Clarion Brookville		
	Pottsville		o e e	A LITERALIE	Succession of sandstone, shales, clays, coals, and thin limestones. Locally, sandstones are open-textured, conglomeratic, massive, cross-bedded, and commonly are found as deposits filling old channels in eroded terrains. Lateral gradations include shale and coal. Non-deposition during the Early Pennsylvanian precluded the development of the basal conglomerate of the Sharon Member in most of southeastern Ohio.	No. No.	3 2 1	Lower Mercer Quakertown Sharon
			Thin, discontinuous zone of impure nodular iron ore and ferruginous sandstone. The unit marks the disconformity between Mississippian and Pennsylvanian strata. Age of the deposit is conjectural, but generally it is included at the base of the Pottsville Formation.					
Mississippian	100000000000000000000000000000000000000	Ondirierentiated	Variable sequence of sandstones and shales; Maxville Limestone is present in patches at the top. In places, various units are conglomeratic and sandstones are massive. Lateral and vertical gradation to siltstone and shale is common. Ground-water potential is limited to extreme western areas. Eastward, the section contains salt water. To the north, post-Mississippian erosion has removed part of the section.					

Figure 6.—Generalized geologic column for southeastern Ohio, including relative position of important coal beds (from Razem and Sedam, 1985).

Fifty-two coal beds are recognized and named in Ohio, most of which are thin and discontinuous. Most minable coals are in the Allegheny and Monongahela Formations, and all are highly volatile and bituminous, falling in the medium (1.1 to 3.0 percent) to high (greater than 3.0 percent) sulfur range (Collins, 1978). Mining has traditionally concentrated on the "numbered" coals—Sharon ("No. 1") through Waynesburg ("No. 11")—which still supply most coal mined in Ohio. The Allegheny and Monongahela Formations are especially productive because of their more uniform thickness and distribution; therefore, most mines in Ohio are in parallel bands following the outcrop of these two formations (Pfaff and others, 1981). The lithologic character of the principal units and the relative positions of the important coals are shown in figure 6.

#### Acknowledgments

The author acknowledges the cooperation of the Ohio Department of Natural Resources, Division of Water; the Muskingum Watershed Conservancy District; the Village of Flushing; and the Village of Lisbon. The author also expresses thanks to area businesses and homeowners in Deersville, Flushing, Lisbon, and Glenford for permitting access to their wells, and to many others throughout eastern Ohio for permitting access to their properties for streamflow measurements.

#### METHODS OF STUDY

One to three long-term surface-water sites in each of the 20 basins (fig. 1) were sampled three times at low flow. A total of 30 sites were sampled twice the first year (1985), and 40 sites were sampled the second year (1986). In each of the five basins selected for "intensive" study, 10 additional surface-water sites were sampled once.

Onsite measurements of discharge, specific conductance, pH, temperature, alkalinity, and acidity were made at each surfacewater site. Discharge was measured by the methods described in Rantz and others (1982). Specific conductance, pH, temperature, and alkalinity were determined by the methods discussed in Skougstad and others (1979). Acidity was determined using the hot-peroxide-treatment method (American Public Health Association, 1975). Water samples for chemical analysis were collected using the equal-transit-rate/equal-width-increment method (U.S. Geological Survey, 1977) for all streams whose water depths were greater than 0.5 foot. These were composited in a churn splitter from which subsamples were drawn. Samples also were sent to the U.S. Geological Survey's National Water Quality Laboratory for analysis of:

Dissolved sulfate;
Total and dissolved iron;
Total and dissolved manganese; and
Total and dissolved aluminum.

Each of the five basins receiving intensive study was examined for the existence of a productive aquifer, that is, an aquifer capable of producing enough water for commercial, light industrial, or public water supplies (excluding unconsolidated deposits along the Ohio River). Where a productive aquifer is present, the ground water was investigated.

At each ground-water sampling site, the following on-site measurements were made:

Water level (when possible); Specific conductance; pH; Temperature; and Alkalinity.

#### Laboratory analysis included:

Total and dissolved aluminum;
Dissolved sulfate;
Total and dissolved iron;
Total and dissolved manganese;
Dissolved silica;
Dissolved calcium;
Dissolved magnesium;
Dissolved sodium;
Dissolved chloride;
Dissolved potassium;
Total organic carbon; and
Total dissolved solids
(residue on evaporation at 180 °C).

#### GENERAL WATER QUALITY

#### Surface Water

The variation in geology throughout Ohio's coal-mining region influences water quality. Basins with limestone bedrock near the surface may buffer acidic mine drainage entering streams, whereas basins without such buffering capacity are more severely affected. Water samples collected from the network of 40 long-term surfacewater sites were distributed near the mouths of major streams in the 20 basins (fig. 1). These sites were sampled twice yearly at low flow. Results of chemical analyses are presented in table 2 (at back of report) and in figure 7.

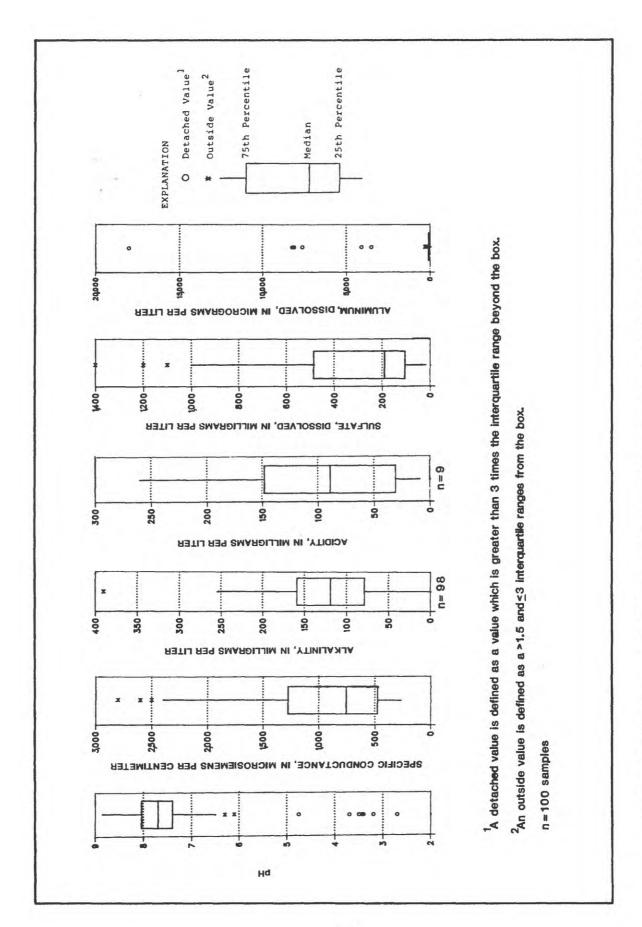


Figure 7.--Box plots showing the range, percentiles, and median values of constituents at the long-term surface-water sites.

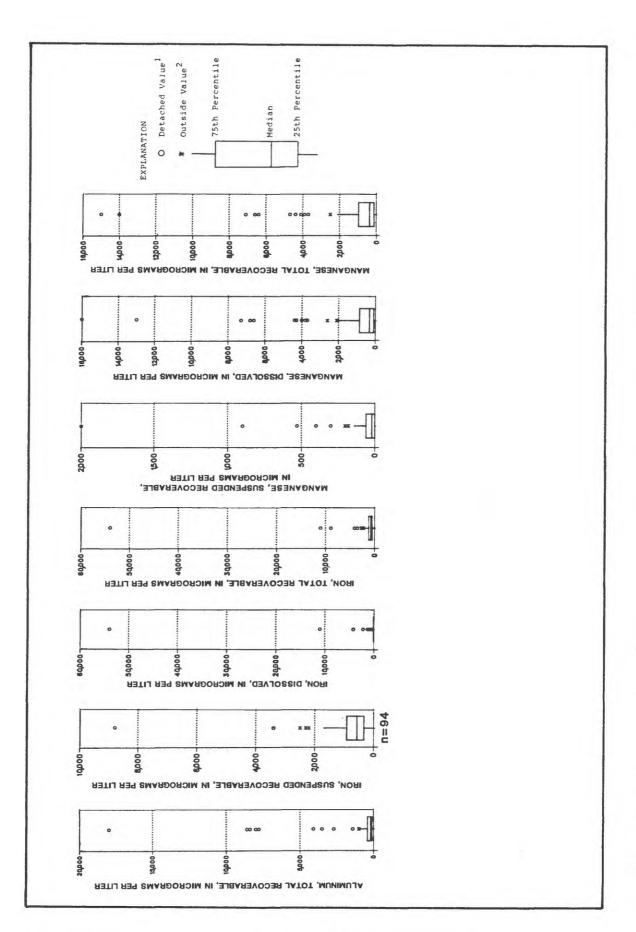


Figure 7.-Box plots showing the range, percentiles, and median values of constituents at the long-term surface-water sites--Continued.

Dissolved-solids concentration is an indicator of general water quality. A rapid method for estimating dissolved-solids concentrations is measurement of specific conductance. Under most conditions, the relation between specific conductance and dissolved-solids concentration is linear. The numerical value for dissolved-solids concentrations, in milligrams per liter (mg/L), generally is 0.55 to 0.75 the value of specific conductance in microsiemens per centimeter at 25  $^{\rm OC}$  (µS/cm) (Hem, 1985). Elevated ionic concentrations in water produce elevated specific-conductance values. In water with elevated dissolved-sulfate concentrations, an elevated measured specific conductance is likely to indicate acid mine drainage. One hundred samples from the long-term site network had a median specific conductance of 760 µS/cm and a range of 270 µS/cm (at site T-1) to 2,800 µS/cm (at site K-1).

The presence of elevated dissolved-sulfate concentrations in surface water can be an indicator of acid mine drainage. Ferrous sulfides in the sandstones and shales overlying coal beds, as well as in waste coal and spoil itself, are important sources of sulfate. Data have shown that, once sulfate concentrations become elevated, they do not return to their premining levels, even after successful reclamation (Pfaff and others, 1981; Hren and others, 1984). Median sulfate concentration for the long-term network was 190 mg/L, with a range of 17 mg/L (at site M-1) to 1,400 mg/L (at site K-1 and O-2). As would be expected, the highest concentrations were found at sites in basins with active or abandoned coal mines.

Pyrite and marcasite are the sulfur-bearing minerals commonly present in coal and associated strata. Pyrite is the most widespread of all sulfide minerals and is recognized as the major source of acid mine drainage in the eastern United States. These minerals are exposed to air and water during coal mining. The oxidation of pyrite may be characterized by the following reactions:

$$FeS_2(s) + 7/2O_2 + H_2O < ---> Fe^{2+} + 2SO4^{2-} + 2H^+$$
 (1)

$$Fe^{2+} + 1/40_2 + H^+ < ---> Fe^{3+} + 1/2 H_2O$$
 (2)

$$Fe^{3+} + 3H_{2}O < ---> Fe(OH)_{3}(s) + 3H^{+}$$
 (3)

$$FeS_2(s) + 14Fe^{3+} + 8H_2O < ---> 15Fe^{2+} + 2SO_4^{2-} + 16H^+$$
 (4)

The oxidation of the sulfide in pyrite to sulfate in reaction 1 releases ferrous iron and acidity (as H<sup>+</sup>) into the water. In reaction 2, ferrous iron is oxidized to ferric iron, which then hydrolyzes to form additional acidity and ferric hydroxide (reaction 3), an insoluble substance often seen in streams receiving acid mine drainage. In reaction 4, sulfide is again oxidized to produce acidity, sulfate, and ferrous iron (Stumm and Morgan, 1970).

When oxygen is present in a mining environment, the rate of acid drainage production is limited by the rate of iron (Fe<sup>+2</sup>) oxidation; the higher the pH, the greater the rate of oxidation (Wiram, 1974). At a pH below 4.5, the rate of oxidation would be extremely slow without microbial catalysis by the iron bacteria Thiobacillus and Ferrobacillus ferrooxidans, which thrive in low-pH water (Stumm and Morgan, 1970). Rock texture, porosity, and permeability also are important in determining the amount of oxidizing agents able to reach iron disulfides, particularly in drift mines and strip-mine highwalls (Pfaff and others, 1981).

Elevated concentrations of iron, manganese, and aluminum are considered undesirable contaminants in terms of aquatic biota and human uses. All three metals are dissolved in surface-water samples from all 20 basins to some degree, and all three are known to be found in increased concentrations during and immediately after mining (Dyer and Curtis, 1977).

Concentrations of iron in water depend on the geology of the area and the chemistry of the water system. More than 1,000 to 2,000  $\mu g/L$  (micrograms per liter) of soluble iron in surface waters generally indicates acidic contamination from mine drainage or other sources (Ward and Wilmoth, 1968). The median concentration of dissolved iron among the long-term sites was 45  $\mu g/L$ , with a range of less than 10  $\mu g/L$  (at sites H-l, L-l, and B-2) to 54,000  $\mu g/L$  (at site Q-2).

Manganese is dissolved from some of the rocks and soils. Small amounts commonly are present in dolomite and limestone and substitute for calcium. Large quantities commonly are associated with elevated iron content and with mine drainage. Manganese is an undesirable impurity in water supplies because of its tendency to deposit black oxide stains. Concentrations of manganese commonly increase after surface mining; manganese commonly is present in concentrations of more than 1,000 µg/L in streams receiving acid mine drainage. Manganese usually stays in solution for greater distances downstream from the pollution source than iron. As the acidity is gradually neutralized, ferric hydroxide precipitates first, followed by manganese (Hem, 1985). Median dissolved-manganese concentration for the long-term network was 320 µg/L, with a range of 10 µg/L (at sites B-l and D-l) to 16,000 µg/L (at site O-2).

Aluminum is present only in negligible quantities in natural waters except in areas where the waters have been in contact with the more soluble rocks of high aluminum content, such as bauxite and certain shales. Acidic waters also commonly contain large amounts of aluminum (Ward and Wilmoth, 1968). Aluminum in public water supplies is not considered a public-health problem under most conditions, as most people ingest 10 to 100 milligrams per day from all sources (McKee and Wolf, 1963). Concentrations of dissolved aluminum in the long-term network ranged from less than 10 µg/L (at sites N-2 and J-1) to 18,000 µg/L (at site Q-1). The median concentration was 30 µg/L.

The pH of natural waters is a measure of acid-base equilibrium achieved by the various dissolved compounds, salts, and gases. The principal system regulating pH in natural waters is the carbonate system, which is composed of carbon dioxide, carbonic acid, bicarbonate ion, and carbonate ion. The pH of samples from the long-term network had a median of 7.70 and ranged from 2.70 (at site Q-2) to 8.85 (at site P-2). In general, the long-term network represented the most downstream point available in each basin. Therefore, a neutral pH value was not uncommon because of possible dilution of any acid mine drainage present. Seven samples had a pH of less than 5.0.

"Buffering capacity," the ability to neutralize additions of acids or bases without a change in pH, is controlled by the concentration of alkalinity and acidity present in water (U.S. Environmental Protection Agency, 1986). Alkalinity resulting from naturally occurring carbonate and bicarbonate is not considered a health hazard in drinking water, and naturally occurring levels up to 400 mg/L as CaCO3 are not considered a problem to human health (National Academy of Sciences, 1974). The median alkalinity in the long-term network was 122 mg/L as CaCO3, with a range of 0 mg/L (at sites Q-1, S-1, and O-1) to 390 mg/L (at site C-2).

Acidity was measured at each site where the pH was 6.5 or less. Above a pH of 6.5, acidity was assumed to be near zero. Nine of the 100 samples had measurable acidity ranging from 9 mg/L as CaCO<sub>3</sub> (at site O-1) to 260 mg/L (at site Q-2), with a median value of 89 mg/L. Elevated acidity is almost always associated with some type of mining activity, as was previously discussed; however, if mining is present where limestone is near the surface, the buffering capabilities of the carbonates may reduce acidity.

The Middle Hocking River, Moxahala Creek, Short and Wheeling Creeks, and Raccoon Creek basins consistently had values at the high end of the range for many constituents, and low pH values. These basins exhibit the characteristics most commonly associated with acid mine drainage. They also were basins with a considerable amount of past as well as present coal mining.

#### Ground Water

Productive aquifers are present in three of the five basins: Stillwater Creek basin, Moxahala Creek basin, and Little Beaver Creek basin. The hydrogeology and results of chemical analyses are presented under the appropriate basin headings. Ground-water samples received supplemental analysis in addition to the constituents analyzed for in the surface-water samples.

Silicon in combination with oxygen, as the oxide SiO2; is termed silica. It is present in solution mainly in water in contact with quartz-bearing rocks. Silica is not readily dissolved in water, although some analyses show increased solubility as water temperature rises (Hem, 1985). Warm ground waters may contain silica concentrations of up to 100 mg/L (Driscoll, 1986). The range of concentrations most commonly observed in natural waters is 1 to 30 mg/L (Hem, 1985). Davis (1964) quotes a median value of 17 mg/L for ground water. Higher concentrations are most commonly related to rock type, rate of movement, water temperature, and natural acids, such as carbonic acid, which affect weathering of rock materials. Although silica does not contribute to the hardness of ground water, it is one of the minerals that forms incrustations in water-supply systems. Silicate scale usually is in the form of calcium or magnesium silicate. Because silicate scale cannot be dissolved by acids or other chemicals commonly used in treatment of wells, water used in boilers must be treated before use by absorption or ion-exchange techniques (Driscoll, 1986). In the three basins having productive aguifers, silica concentrations ranged from 8.7 to 19 mg/L.

Hardness of water is defined as its content of metallic ions that react with sodium soaps to produce scummy residue and that react with negative ions, when the water is evaporated in boilers, to produce solid boiler scale (Camp, 1963). In freshwater, these ions are primarily calcium and magnesium. Hardness commonly is expressed as an equivalent concentration of calcium carbonate (CaCO<sub>3</sub>). The following is a commonly used classification (U.S. Environmental Protection Agency, 1986):

Classification of Water by Hardness Content

[Concentration in mg/L as CaCO3]

0-75	soft
75-150	moderately hard
150-300	hard
300 and up	very hard

Hardness is derived principally from calcium and magnesium. Its natural source is limestone dissolved by slightly acidic rainwater in contact with carbon dioxide. Industrial sources may include abandoned and active mines. Hardness is commonly reported in carbonate and noncarbonate fractions. The carbonate fraction is chemically equivalent to the bicarbonates present. Bicarbonates generally are measured as alkalinity; therefore, carbonate hardness usually is considered equal to the alkalinity (U.S. Environmental Protection Agency, 1986). If the hardness exceeds the alkalinity, the excess is reported as "noncarbonate hardness" (Hem, 1985). In the three basins having productive aguifers, hardness ranged from 100 to 840 mg/L.

Calcium is a major constituent of many common rock minerals. The most common forms in sedimentary rocks are carbonates. In sandstone, calcium carbonate commonly is present as a cement between particles. It is an essential element for plant and animal life forms and is a major component of the solutes in most natural water (Hem, 1985). In the three basins having productive aquifers, calcium concentrations ranged from 27 to 250 mg/L.

Magnesium is present in significant amounts in most limestones and dolomites. Dissolution of these materials brings magnesium into solution, but the process is not easily reversed. Precipitate from this solution may be nearly pure calcite. Concentrations of magnesium tend to increase along a ground-water flow path in limestones undergoing dissolution until a relatively high calcium:magnesium ratio is obtained (Hem, 1985). In the three basins having productive aquifers, magnesium concentrations ranged from 8.3 to 64 mg/L.

Sodium may be present in unaltered mineral grains as part of the cementing material or as crystals of sodium salts deposited with sediments left by saline water. After dissolution, sodium tends to remain dissolved. There are no important precipitation reactions that can maintain low sodium concentrations in water similar to the way that carbonate precipitation controls calcium concentration (Hem, 1985). The U.S. Environmental Protection Agency (1986) recommends a limit of 20 mg/L of sodium in water for very restricted sodium diets and 270 mg/L for moderately restricted diets. No specific criterion is set for drinkingwater supplies. In the three basins having productive aquifers, sodium concentrations ranged from 4.6 to 73 mg/L.

Potassium is slightly less common than sodium in igneous rock but more abundant in all the sedimentary rocks. In most natural water, potassium concentrations are much lower than sodium concentrations. Potassium is commonly present in clay minerals and some evaporite rocks. Concentrations of potassium more than a few tens of milligrams per liter are very unusual except for waters with high dissolved-solids concentration (Hem, 1985). In the three basins having productive aquifers, potassium concentrations ranged from 1.1 to 4.5 mg/L.

Total organic carbon (TOC) is widely used as a measure of the total concentration of organic constituents in water. The determination of total organic carbon gives a better measure of organic matter present in aqueous solution and (61) suspension than does chemical oxygen demand (Goerlitz and Brown, 1972). In ground wa'er, total organic carbon is likely to be equal to dissolved organic carbon. Thurman (1985) lists 0.7 mg/L as an approximate concentration for total organic carbon typical of natural ground water, most all of which occurs in the dissolved state. In the three basins having productive aquifers, concentrations of total organic carbon ranged from 0.4 to 6.0 mg/L.

Excessive dissolved solids are objectionable in drinking water for aesthetic reasons. Dissolved solids consist of inorganic salts, small amounts of organic matter, and dissolved materials (U.S. Environmental Protection Agency, 1986). For drinking water, the Ohio Environmental Protection Agency (1978) limits dissolved solids to 500 mg/L. In the three basins having productive aquifers, dissolved-solids concentrations ranged from 194 to 1,200 mg/L.

Because many ground-water sources are domestic or public supplies, the following Ohio Environmental Protection Agency (1978) water-quality standards are listed for reference:

Chlorides	250	mg/L
Dissolved solids	500	mg/L
Iron (dissolved)	300	µg/L
Manganese	50	µg/L
Sulfates		mq/L

All concentrations are expressed as total concentrations unless otherwise noted.

GEOLOGIC SETTING AND WATER QUALITY OF SELECTED BASINS

#### Stillwater Creek Basin

#### Geologic Setting

The Stillwater Creek basin (fig. 2) is within the stream-dissected Appalachian Plateaus physiographic province (Fenneman, 1938). The landscape is one of relatively narrow to broad stream valleys. Most of the basin is thoroughly dissected by drainage-ways that have cut deep valleys and left narrow ridgetops.

Bedrock beneath the Stillwater Creek basin consists of alternating layers of sandstone, shale, limestone, clay, and coal of Pennsylvanian age. The rock strata crop out in a westto-east pattern from older to younger formations. The Allegheny Formation underlies the northwestern part of the basin and consists of mostly sandstones, some shales, and coal. Most of the basin is covered by strata of the Conemaugh Formation, which consist of shales, sandstones, limestones, and coal. In the extreme eastern part of the basin is a relatively small outcropping of the Monongahela Formation, which is made up of shales, limestones, sandstones, and some coal. The sandstone formations to the north in the basin may yield as much as 25 gal/min (gallons per minute). These sandstones are not uniform in thickness, character, or areal extent; therefore, locally large yields do not indicate extensive water-bearing sandstone deposits. Except for localized areas of thick sandstones, bedrock in the northern quarter of the basin will supply enough water for domestic and farm uses only (Walker, 1962a).

Logs on file at the Ohio Department of Natural Resources show that beneath the floodplains of parts of Stillwater and Little Stillwater Creeks are deposits of sand and gravel interbedded with silt and clay that may yield as much as 50 gal/min. The materials filling these valleys were deposited from glacial meltwaters to the north (Walker, 1962a). Continued exploration in these valleys would be needed to determine the extent of these permeable, water-bearing sands and gravels.

The fill elsewhere along Stillwater and Little Stillwater Creeks and their tributaries ranges from 20 to 200 feet in thickness, and consists largely of clay, silt, and fine sand, which is a poor source of ground water. Water-supply wells normally are drilled through these deposits to the underlying bedrock, although large-diameter dug wells may be developed. Similar materials are present in most stream valleys throughout the basin. Rarely, sand and gravel deposits of limited extent are encountered that can yield 25 gal/min or more. Such deposits have been located along Jockey Hollow Run north of Flushing. However, these are unusual and not found in most areas (Walker, 1962a).

In most of the basin, sandstone and sandy shale bedrock will supply limited quantities of water for domestic and farm use. Yields commonly average less than 3 gal/min from drilled wells and may be supplemented by springs or dug wells. Wells drilled to depths greater than 200 feet below stream level are likely to encounter saline water (Walker, 1962a).

Ground-water levels in the Stillwater Creek basin (fig. 8) were measured near the Village of Flushing and at Tappan Lake Park near Deersville. Water levels are presented in table 3.

The Village of Flushing water supply is located along the outwash valley of Jockey Hollow Run northwest of the village (fig. 9). The entire area is surrounded by land which has been or is being surface mined. Well records from the Village of Flushing indicate unusually high-yielding sand and gravel deposits with a developed capacity of up to 200 gal/min. There were not enough wells in the area to construct a contour map of the potentiometric surface from the water-level measurements.

Tappan Lake Park has drilled a total of 16 wells into the underlying shale and sandstone of the Conemaugh Formation as water supplies for recreational use. Five of these wells have been abandoned. Two of the wells have been allocated for use to the Village of Deersville. Of the remaining wells, one well, HR-32, supplies the cabin area, and the water is untreated. Water from wells HR-30, HR-33, HR-40, HR-41, and HR-43 is combined and processed through a treatment plant and sent to a holding tank before distribution to the park. Thirteen water-level measurements were obtained. Well locations are shown in figure 10.

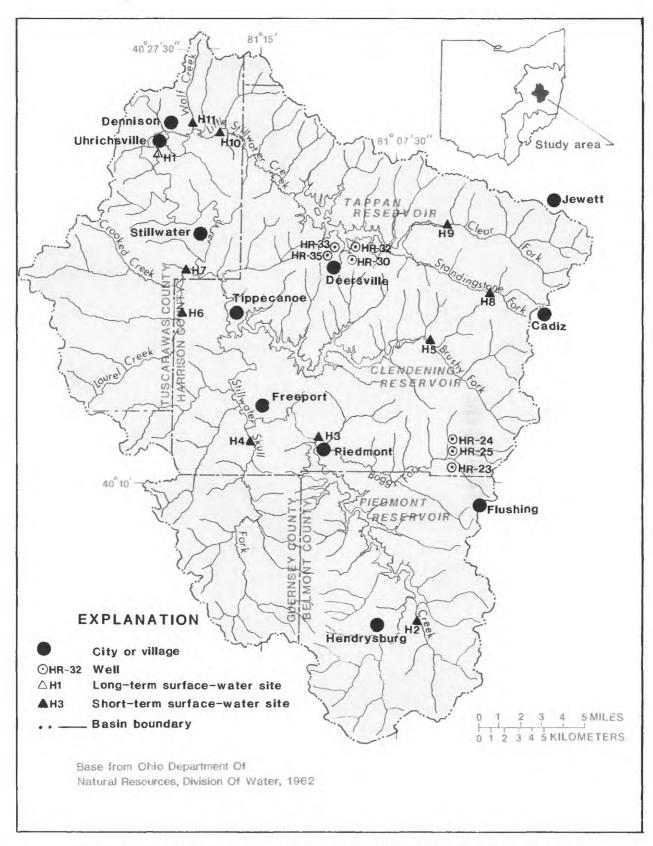


Figure 8.--Short-term and long-term surface-water sites and ground-water sites in the Stillwater Creek basin.

Table 3.--Ground-water-level measurements

Local number	Latitude	Longitude	Date	Water level
	Little Beaver	Creek basin	(Columbiana County)	
CO-1	404754	804537	09-02-86	44.87
CO-2	405008	804608	09-03-86	flowing
CO-3	404639	804833	09-02-86	44.20
CO-4	404846	804610	09-03-86	23.00
CO-5	404553	804317	09-02-86	28.89
CO-6	404506	804141	09-02-86	47.87
CO-7	404613	804341	09-03-86	107.00
CO-8	404456	803912	09-04-86	133.62 8.55
CO-9 CO-10	404545 404620	804149 804151	09-03-86 09-03-86	8.73
CO-10	404843	804523	09-03-86	35.57
CO-11	404940	804557	09-04-86	33.74
CO-12	404906	804513	09-04-86	40.59
CO-14	404847	804550	09-04-86	34.51
CO-15	404841	804650	09-04-86	98.23
CO-16	404751	804637	09-03-86	18.85
CO-17	404716	804548	09-03-86	44.87
CO-18	404546	804527	09-03-86	34.16
CO-19	404607	804625	09-03-86	5.28
CO-20	404705	804805	09-04-86	13.84
CO-21	404644	804839	09-03-86	74.39
CO-22	404652	805002	09-04-86	24.60
CO-23	404843	804608	09-01-86	23.00
	Stillwater	Creek basin	(Harrison County)	
HR-23	401043	810528	11-21-85	
HR-24	401049	810547	11-21-85	
HR-25	401047	810547	11-21-85	16.80
HR-26	401045	810547	11-21-85	11.70
HR-27	401028	810514	11-21-85	flowing
HR-28	401011	810459	11-21-85	4.00
HR-29	401045	810533	11-21-85	3.00
HR-30	401920	811043	11-22-85	9.35
HR-31	401907	811045	11-22-85	9.60
HR-32	401937	811030	11-22-85	34.00
HR-33	401917	811112	11-22-85	10.00
HR-34	401915 401858	811117 811122	11-22-85 11-22-85	10.50
HR-35 HR-36	401903	811111	11-22-85	27.80
111720			11-22-85	26.50
HR-37 401856 811123 HR-38 401859 811117			11-22-85	2h - 511

Table 3.--Ground-water-level measurements--Continued

Local number	Latitude	Longitude	Date	Water level <sup>l</sup>
HR-39	401856	811115	11-22-85	44.00
HR-40	401908	811111	11-22-85	18.00
HR-41	401911	811107	11-22-85	10.20
HR-42	401917	811039	11-22-85	14.80
HR-43	401912	811104	11-22-85	10.70
	Moxahala	Creek basin	(Perry County)	
PE-39	395204	821714	08-19-86	44.40
PE-40	395457	822431	08-18-86	3.75
PE-42	395336	821406	08-18-86	35.20
PE-43	395316	821902	08-19-86	18.78
PE-44	395320	821906	08-18-86	36.87
PE-45	395309	821925	08-19-86	
PE-46	395353	821914	08-19-86	22.69
PE-47	395507	822105	08-18-86	0.0
PE-48	395424	822146	08-18-86	26.89
PE-49	395218	821517	08-22-86	1.08
PE-50	395431	821955	08-22-86	
PE-51	395308	821450	08-18-86	9.68
PE-52	395346	821920	08-18-86	21.02
PE-53	395318	821826	08-18-86	74.80
PE-54	395428	822129	08-18-86	18.18
PE-55	395507	822150	08-18-86	2.40
PE-56	395401	821908	08-18-86	36.87

<sup>&</sup>lt;sup>1</sup>Below land surface.

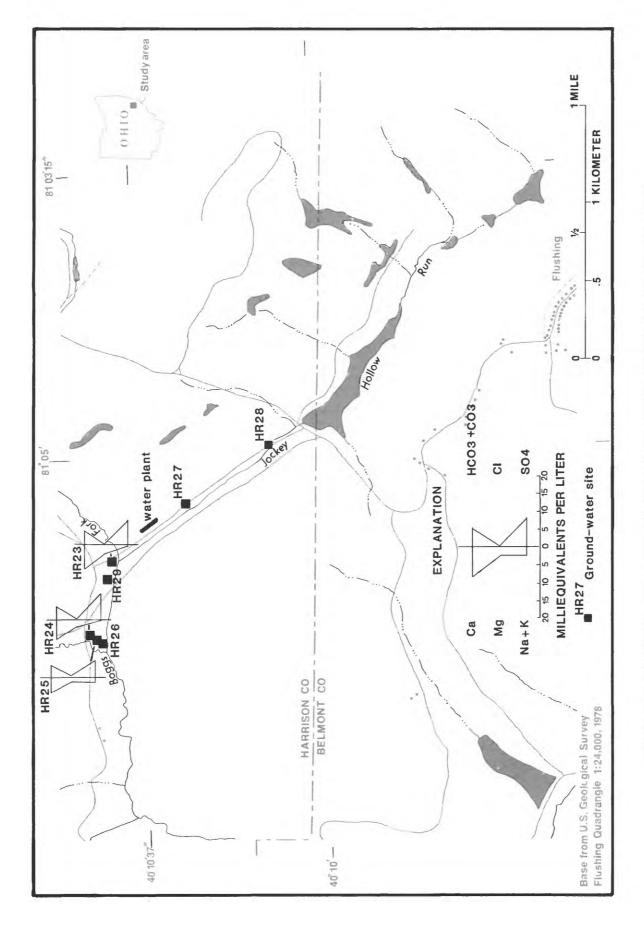


Figure 9.--Ground-water sites and Stiff diagrams for Stillwater Creek basin, Flushing, Ohio.

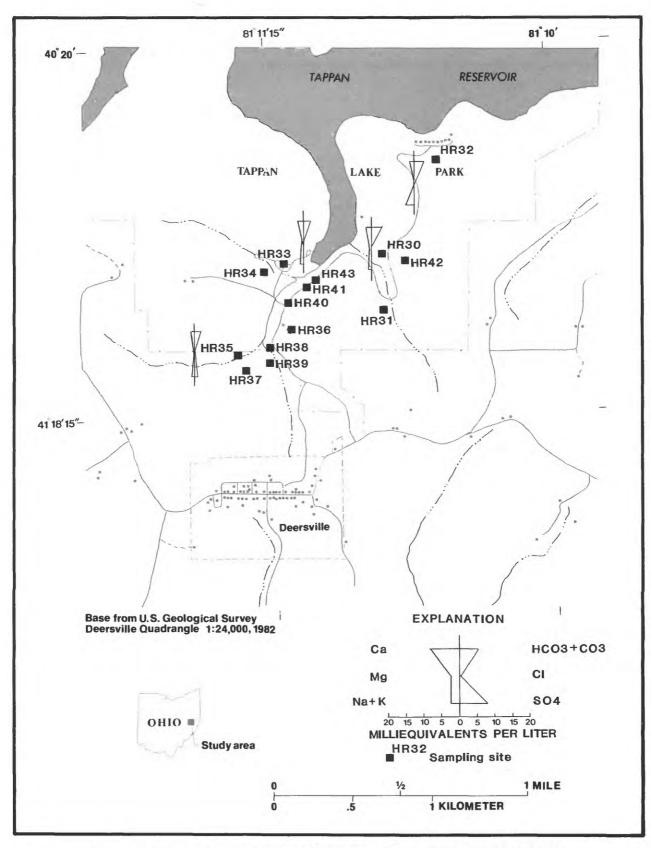


Figure 10.—Ground-water sites and Stiff diagrams for Stillwater Creek basin,
Tappan Lake Park, Deersville, Ohio.

On the basis of the limited water-level data available, the ground water at Tappan Park appears to be flowing toward Tappan Lake. The wells closest to the lake approach or nearly match the level of the lake.

#### Water Quality

Surface-water quality of the Stillwater Creek basin was investigated by sampling ll stream sites at low flow (fig. 8). Ground-water quality was investigated by sampling seven sites in two distinct aquifers in the basin, at the Village of Flushing, Ohio, and at Tappan Lake Park near Deersville, Ohio (figs. 9 and 10).

#### Surface Water

A total of 11 water-quality samples were collected from Stillwater Creek (2 sites), Boggs Fork, Brushy Fork, Clear Fork, Crooked Creek, Little Stillwater Creek, Laurel Creek, Skull Fork, Standingstone Fork, and Wolf Run, during the period September 30 through October 2, 1985 (fig. 8). Results of chemical analyses are presented in table 4 and figure 11.

Specific conductance ranged from 390  $\mu$ S/cm (at site H-6) to 3,300  $\mu$ S/cm (at site H-5), with a median of 1,300  $\mu$ S/cm.

Values for pH ranged from 7.00 (at site H-4) to 8.00 (at site H-5), with a median of 7.60. Alkalinity concentrations ranged from 78~mg/L as  $\text{CaCO}_3$  (at site H-4) to 372~mg/L (at site H-11), with a median of 212 mg/L.

The dissolved-sulfate concentration ranged from 15 mg/L (at site H-6) to 1,900 mg/L (at site H-5), with a median of 580 mg/L.

Total-iron concentrations ranged from 290  $\mu g/L$  (at site H-9) to 16,000  $\mu g/L$  (at site H-10), with a median of 850  $\mu g/L$ . Dissolved-iron concentrations ranged from less than 10  $\mu g/L$  (at site H-1) to 130  $\mu g/L$  (at site H-3), with a median of 30  $\mu g/L$ .

Concentrations of total aluminum ranged from 100  $\mu$ g/L (at sites H-7 and H-9) to 6,500  $\mu$ g/L (at site H-10), with a median of 300  $\mu$ g/L. Dissolved-aluminum concentrations ranged from less than 100  $\mu$ g/L (at sites H-5, H-7, H-8, and H-9) to 300  $\mu$ g/L (at sites H-1 and H-3), with a median of 100  $\mu$ g/L.

Total-manganese concentrations ranged from 160  $\mu$ g/L (at site H-2) to 2,900  $\mu$ g/L (at site H-10), with a median of 310  $\mu$ g/L. Dissolved-manganese concentrations ranged from 90  $\mu$ g/L (at site H-5) to 1,700  $\mu$ g/L (at site H-10), with a median of 220  $\mu$ g/L.

#### Ground Water

Three ground-water samples were obtained from the sand and gravel deposits of the water supply at the Village of Flushing (fig. 9), and four samples were obtained from shallow bedrock wells (HR-30, HR-32, HR-33, and HR-35) at Tappan Lake Park near Deersville (fig. 10). A corresponding Stiff (1951) water-quality diagram also is shown for each sampled well. Table 5 lists the results of chemical analyses. Stiff water-quality diagrams are shown in figure 10.

Most of the ground water in the Flushing area (fig. 9) can be classified as a very hard calcium sulfate type that is high in dissolved solids. In comparison to other outwash aquifers in Ohio (Evans, 1977; deRoche and Razem, 1984; Norris and Fidler, 1969), the water at Flushing appears to have elevated values of specific conductance (1,100 to 1,350 µS/cm) and total dissolved solids (883 to 1,200 mg/L). Values for pH ranged from 7.10 to 7.50, which is typical. The high carbonate concentration in water in the Conemaugh Formation may be helping to buffer acids formed as a result of mining activities if water from the bedrock below were flowing with an upward vertical component. Alkalinity ranged from 274 to 340 mg/L as CaCO<sub>3</sub>.

Calcium is the most abundant cation, and ranges from 170 to 250 mg/L for the three samples. Sulfate, the most abundant anion, ranged from 390 to 590 mg/L, which exceeds the Ohio Environmental Protection Agency (OEPA) standard for public supply and is much greater than the sulfate concentration of water from other outwash aquifers. Concentrations of magnesium ranged from 31 to 54 mg/L, and sodium concentrations ranged from 20 to 56 mg/L, both slightly elevated compared with other outwash aquifers in Ohio. Hardness ranged from 550 to 840 mg/L as CaCO<sub>3</sub>. Potassium concentrations ranged from 1.1 to 1.4 mg/L, which is typical of outwash aquifers in the State. Concentrations of silica and chloride ranged from 10 to 12 mg/L and 3.7 to 15 mg/L, respectively.

Concentrations of dissolved iron ranged from 17 to 3,300 µg/L; concentrations in two samples exceeded the OEPA standard for public supply. The lowest value (17 µg/L) was found in HR-24, where most of the dissolved iron and manganese had precipitated out after filtration. The precipitation was possibly caused by inadvertent aeration of the water before the sample could be taken. Total-iron concentrations ranged from 1,400 to 3,500 µg/L. Total concentrations of manganese ranged from 1,400 to 3,100 µg/L and dissolved manganese from 780 to 3,100 µg/L. All three of the samples exceeded the OEPA standard. Concentrations of total aluminum were 100 µg/L in all three samples. Total organic carbon ranged from 2.2 to 4.2 mg/L, which is considerably higher than the approximate natural concentration of 0.7 mg/L for Ohio ground water reported by Thurman (1985).

[deg. C, degrees Celsius; ft<sup>3</sup>/s, cubic feet per second; mg/L, milligrams per liter;

Date	Stream- flow, instan- tangous (ft /s)	du ta	fic	ph (stand- ard units)	Temper- ature (deg. C)	Acidity (mg/L as CaCO <sub>3</sub> )	Alka- linity, field (mg/L as CaCO <sub>3</sub> )	Sulfate dis- solved (mg/L as SO <sub>4</sub> )	Alum- inum, total recov- erable (µg/L as Al)
03127500		H-1 S							081 20 50w)
Sep 1985 30	12	1	,200	7.60	19.0		206	390	700
400506081	073900	H-2	Still	water C	nr Hendry	sburg OH	(lat 40 0	5 06n long	081 07 39w
Oct 1985 02					14.0			1,300	300
401126081	121600	H-3	Boggs	Fork at	Piedmont	OH (lat	40 11 26r	long 081	
Oct 1985								910	
03126170		H-4	Skul1	F at Fr	eeport OH	(lat 40	11 52n 1	ong 081 16	i 13w)
Sep 1985 30					14.0		78		300
								ong 081 07	01w)
Oct 1985 01	3.7	3,	,300	8.00	14.0		285	1,900	
401647081	194200	H-6	Laure	l Cr nr	Tippecanoe	OH (lat	40 16 47	n long 081	
Sep 1985 30					18.0		155		400
03127100		H-7	Crook	ed C nr	Stillwater	OH (lat	40 18 29	n long 081	19 26w)
Sep 1985 30				7.40			212	19	100
01724081	032100	H-8	Stand	ingstone	F nr Cadi	z OH (la	t 40 17 1	6n long 08	1 02 33w)
Oct 1985 01								1,400	
402012081	051300	H-9	Clear	Fork nr	Jewett OH	(lat 40	20 12n 1	ong 081 05	13w)
Oct 1985					14.5			910	
3128600					C nr Denn		(lat 40 2	4 19n long	081 17 18w)
Oct 1985 01	8.4		710	7.20	15.0		177	230	6,500
02429081	185300	H-11	Wolf	Run nr		H (lat 4		long 081 1	8 53w)
Oct 1985 01	0.8	1,	300	7.90	11.5		372	340	300

#### surface-water sites in the Stillwater Creek basin

Alum-			235			Manga-	
inum,	e en		Iron,		Manga-	nese,	
	lum-	Iron,	sus-		nese,	sus-	Manga-
	num,		pended	Iron,	total	pended	nese,
	dis-		recov-	dis-	recov-	recov-	dis-
erable s	olved	erable		solved	erable	erable	solved
	μg/L)	(µg/L)	-	(µg/L)	(µg/L)		
as Al) a	s Al)	as Fe)	as Fe)	as Fe)	as Mn)	as Mn)	as Mn)
03127500	H-1	Stillwate	r C at Ul	richsvil	e OH (lat	40 23 10	n long 081 20
400	300	1,700		<10	1,300	900	400
400506081073900	H-2	Still wate	r C nr He	ndrysburg	OH (lat	40 05 06n	long 081 07 3
200	100	580	530	50	160	20	140
401126081121600			k at Pied	mont OH	lat 40 11	26n long	081 12 16w)
0	300	380	250			110	200
03126170	H-4	Skull F a					1 16 13w)
200	100	540	490	50	530	120	410
401538081070100	н-5	Brushy Fo	rk nr Cad	liz OH (la	t 40 15 3	8n long 0	81 07 01w)
	<100	1,200	1,200	30	210	120	90
401647081194200							g 081 19 42w)
300	100	1,100	1,000	60	640	110	530
03127100				water OH		8 29n lon	g 081 19 26w)
	<100	850	760	90	720		680
401724081032100		Standings					
	<100	530	520	10	220	10	210
402012081051300	н-9	Clear For	k nr Jewe	tt OH (la	t 40 20 1	2n long 0	B1 05 13w)
	<100	290	270		250	30	220
03128600	H-10	L Stillw	ater C nr	Dennison	OH (lat	40 24 19n	long 081 17 16
6400	100	16,000	16,000	10	2,900	1,200	1,700
402429081185300	H-11	Wolf Run	nr Denni	son OH (1	at 40 24	29n long	081 18 53w)

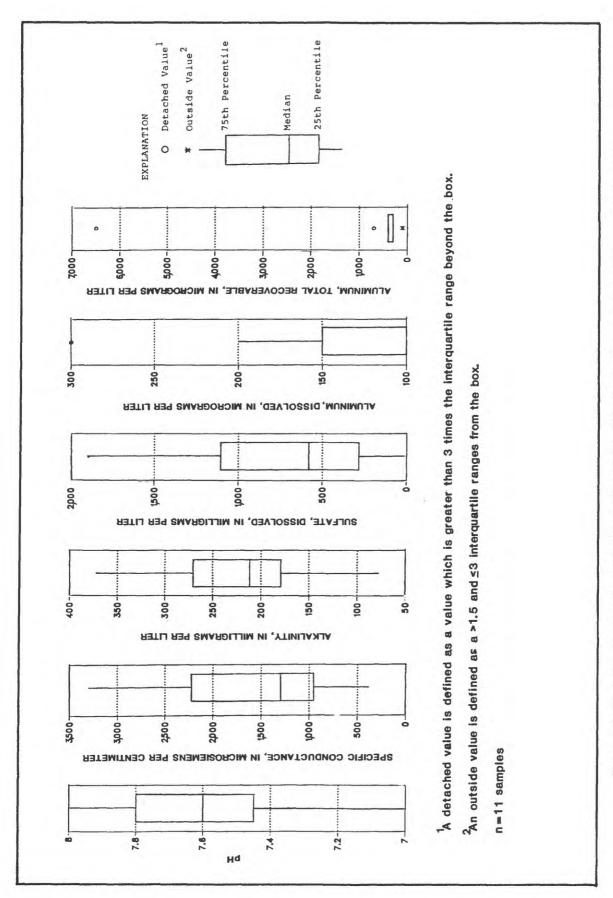


Figure 11.--Box plots showing the range, percentiles, and median values of constituents at surface-water sites in the Stillwater Creek basin.

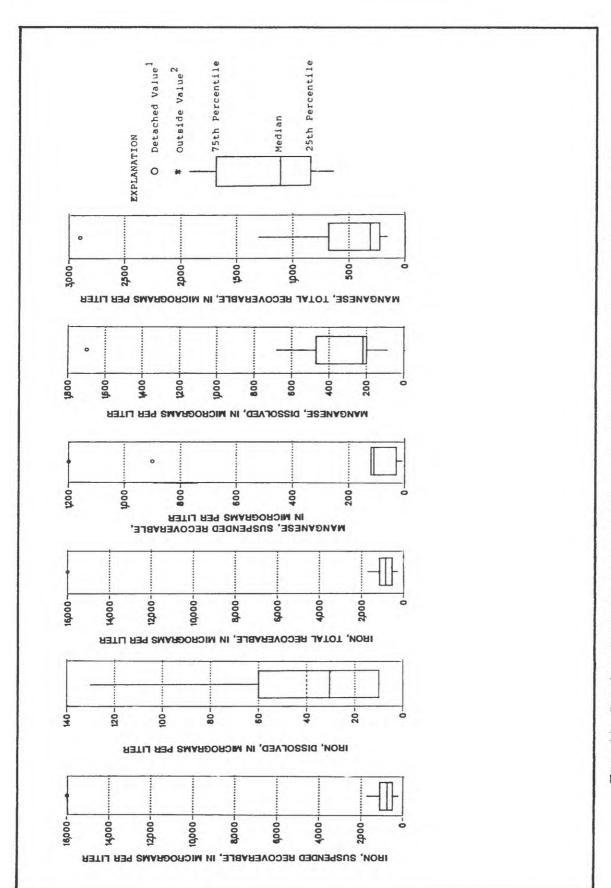


Figure 11.--Box plots showing the range, percentiles, and median values of constituents at surface-water sites in the Stillwater Creek basin--Continued.

Tappan Lake Park (fig. 10), the second area in the Stillwater Creek basin to be evaluated for ground-water quality, is located near Deersville, and is owned and administrated by the Muskingum Watershed Conservancy District. No coal has been mined in this area. The results of this evaluation will provide background data on the quality of shallow bedrock aquifers in this part of Ohio.

Analyses show that Tappan Park has a moderately hard to hard calcium bicarbonate type ground water. Dissolved-solids concentrations ranged from 194 to 311 mg/L, which is below the limit of 500 mg/L recommended by the U.S. Environmental Protection Agency (1986) for a public supply and is slightly below the average noted in other bedrock aquifers in Ohio (Razem and Sedam, 1985). Specific conductance ranged from 320 to 500  $\mu\text{S/cm}$ . Values for pH ranged from 7.20 to 7.70. Alkalinity ranged from 173 to 263 mg/L as CaCO3.

Calcium is the most abundant cation with a range of 27 to 50 mg/L for the four samples. One sample (HR-35) did not have a dominant cation, and one sample (HR-32) had unusually high concentrations of sodium and potassium (73 mg/L and 3.2 mg/L, respectively). Bicarbonate, the most abundant anion, was present in concentrations of 211 to 321 mg/L. Magnesium concentrations ranged from 8.3 to 12 mg/L, and sodium concentrations ranged from 21 to 73 mg/L. Hardness ranged from 100 to 170 mg/L as CaCO<sub>3</sub>, which is classified by the USEPA as moderately hard to hard. Potassium concentrations ranged from 1.5 to 3.2 mg/L. Silica and chloride concentrations ranged from 11 to 17 mg/L and from 3.7 to 14 mg/L, respectively. Total organic carbon in the Tappan wells ranged from 0.8 to 1.4 mg/L.

Concentrations of iron (1,400 to 2,000  $\mu g/L$  total and 710 to 1,200  $\mu g/L$  dissolved) and manganese (80 to 370  $\mu g/L$  total) exceeded the OEPA standards for public supply in the four samples. Total aluminum was 100  $\mu g/L$  in three samples; in water from HR-30, the total aluminum concentration was 600  $\mu g/L$ .

Table 5.--Ground-water analyses for the Stillwater Creek basin, Flushing, Ohio, and Tappan Lake Park, Deersville, Ohio

[deg. C, degrees Celsius; mg/L, milligrams per liter; µg/L, micrograms per liter; µS/cm, microsiemens per centimeter at 25 degrees Celsius]

Date	below land surface (water level) (feet)	Spe- cific con- duc- tance (µS/cm)	ph (stan- dard units)	Temper- ature (deg. C)	Oxygen, dis- solved (mg/L)	Hard- ness (mg/L as CaCO <sub>3</sub> )	ness, noncar- bonate, field (mg/L as	Calcium dis- solved (mg/L as Ca)	Magne- sium, dis- solved (mg/L
Flushing wel	vells:								
401043081052	)52800 HR-2	3 (	40 10 43n	1 -	05 2				
11/21/85		1,350	7.10	10.0	2.0	840	200	250	53
401049081054	054700 HR-2	-24 (lat	40 10 49n	n long 081	05 4			 	1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1
11/21/85		1,350	7.30	10.0	1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	770	460	220	54
401047081054700	HR	-25 (lat	40 10 47n	1 -	05 47w)		1		1 1 1 1
11/21/85	16.80	1,100	7.50	10.0		550	280	170	31
Tappan wells	115:								
401920081104300	HR	-30 (lat	40 19 20n	10	10 43w)				
11/22/85		400	7.40	11.0	0.7	170	0	50	12
401937081103000	HR	-32 (lat	40 19	lon	. 10 30w)				1
11/22/85 34.00	34.00	200	7.65	1	0.4	100	0	27	8
40191708111	11200 HR		40 19	lon	11 12		1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1		
11/22/85	10.00	380	10		1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	150	0	42	10
40185808111	2200 HR	-35 (lat	40 18 58n	in long 081	11 22w)		1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1		1
11/22/85		320	7.20	11.5	0.2	120	0	33	9.4

	Sodium	н	Bicar	Alka-		Chlo-	Silica,	Solids, residue at 180
Date	dis- solved (mq/L	dis- solved (mq/L	bonate, field	linity, field (mq/L as		dis- solved	solved (mg/L	deg. C dis- solved
	as Na)	as K)	HCO <sub>3</sub> )	caco <sub>3</sub> )	as SO <sub>4</sub> )	as Cl)	SiO <sub>2</sub> )	(mg/L)
Flushing wel	wells:							
401043081052800	052800	HR-23 (lat	40 10	43n long 0	081 05 28w)		1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	
11/21/85	20	1.4	415	340	260	6.2	12	1,200
401049081054	1054700	HR-24 (lat	40 10	ng.	081 05 47w)	1		1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1
11/21/85	20	1.4	371	304	590	3.7	12	1,200
401047081054		HR-25 (lat	40 10	47n long 0	081 05 47w)			
11/21/85	56	1.1	334	274	390	15	10	883
appan	wells:							
401920081104		HR-30 (lat	40 19	20n long 0	081 10 43w)			
11/22/85	21	1.7	306	251	18	6.0	14	248
401937081103	1103000	HR-32 (lat	40 19	37n long 0	081 10 30w)	 	t t t 1	
11/22/85	73	3.2	321	263	11	14	11	311
401917081111	1111200	HR-33 (lat	40 19	1 6	081 11 12w)			
11/22/85	2	1.5	257	211	6.6	3.7	16	222
401858081112	1112200	HR-35 (lat	40 18	58n long 0	081 11 22w)			
11/22/85	22	2.3	211	173	12	6.1	17	194

				Al um	1					
	Solids,			inum,		6 5		Manga-	3	
	constit	total	minum.	pended	ded	total	Tron.	total	manga-	Carbon.
	uents.		dis-		- 00	recov-	dis-	recov-	dis-	organic
	dis-	erable	solved		ble	erable	solved	erable	solved	total
Date	solved (mg/L)	l (µg/L as Al)	(µg/L as Al)	as as	/L A1)	(ng/L as Fe)	(ng/L as Fe)	(ug/L as Mn)	(µg/L as Mn)	(mg/L as C)
Flushing wells:	wells:									
401043081052800 HR-23	1052800	HR-23 (lat	40 10	43n long	9 081	05 28w)	1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1			
11/21/85	1,100	100	<100			3,500	3,300	3,100	3,100	2.2
401049081054700	1054700	HR-24 (lat	40 10	49n long	9 081	05 47w)				
11/21/85	1,100	100	<100			3,000	17	2,800	780	2.7
401047081054700	1054700	HR-25 (lat	40 10	47n long	9 081	05 47w)				
11/21/85	840	100	100		0	1,400	1,100	1,400	1,600	4.2
Tappan wells:	ells:									
401920081104300 HR-30	1104300	HR-30 (lat	40 19	20n long	9 081	10 43w)				
11/22/85	270	009	<100			1,400	710	250	300	1.4
401937081103000	1103000	HR-32 (lat	40 19	37n long		081 10 30w)				
11/22/85	310	100	100		0	1,800	1,200	80	72	0.8
401917081111200	1111200	HR-33 (lat	40 19	17n long	1	081 11 12w)				
11/22/85	230		<100			1,800	730	370	390	1.4
401858081112200	1112200	HR-35 (lat	40 18	58n long		081 11 22w)				
			1							

-

## Symmes, Ice, and Indian Guyan Creeks Basin

## Geologic Setting

The Symmes, Ice, and Indian Guyan Creeks basin (fig. 2) is principally in the Lexington Peneplain section of the Appalachian Plateaus physiographic province (Fenneman, 1938). It is deeply dissected by streams and is rugged and hilly, with numerous, V-shaped steep valleys with steeply sloping sides. Most streams are in a youthful stage -- that is, they have no flood plains and are actively downcutting. Symmes Creek, however, is in a mature stage, and the valley is broad and bordered by low, rounded hills. The western part of the basin, including Ice Creek, is largely underlain by sandstone of the Allegheny Formation, which forms steep cliffs or outlying knobs or hills in many areas. Commonly, streams have steep, rock-walled, overhanging valleys. The Symmes and Indian Guyan Creeks valleys to the east are underlain by more shaly rock of the Conemaugh and Monongahela Formations, which are characterized by more gentle slopes. The divides in the eastern part of the basin are long and broad.

The subsurface of the Symmes, Ice, and Indian Guyan Creeks basin is composed of layers of shale, sandstone, coal, and limestone of Pennsylvanian age. Because of the slight dip of the rock strata to the east at an average of 25 feet per mile (Stout, 1916), these layers crop out in irregular north-south bands across the basin and are successively younger from west to east (Maxey, 1940). The Allegheny Formation, which consists of mostly sandstone and shale, crops out in the extreme western and north-central parts of the basin. The Conemaugh Formation, which consists of shales, limestone, coal, and sandstone, covers the central part of the basin in a wide belt. Only in the extreme eastern part of the basin does the Monongahela Formation crop out. It consists of shales, limestones, coal, and sandstones (Pree, 1962).

Pleistocene and Holocene alluvial deposits of gravel, sand, silt, and clay overlie the bedrock throughout the uplands and generally are less than 50 feet thick. The flood plains of many of the streams in the basin are underlain by thin deposits composed of clay, silt, and only small amounts of sand and gravel. The deposits along the Ohio River, which consist of thick, permeable sand and gravel and interbedded clays, are more than 78 feet thick along the reach of river in this basin and some wells yield as much as 1,000 gal/min with proper development (Pree, 1962).

The bedrock in this basin yields generally less than 5 gal/min. Most wells are finished in sandstone of the Conemaugh Formation. Logs from these wells indicate that the sandstone beds generally are thin and range from 5 to 30 feet in thickness. In many parts of the basin, cisterns are used as a supplementary source of water (Pree, 1962).

All streams in the basin flow into the Ohio River. A few streams in the extreme eastern and southern parts of the basin that drain directly into the Ohio River (that is, are not tributary to the Symmes, Ice, or Indian Guyan Creeks) are included in this basin for purposes of this study.

## Water Quality

Surface-water quality of the Symmes, Ice, and Indian Guyan Creeks basin was investigated by sampling 12 stream sites at low flow (fig. 12). Ground-water quality was not investigated in this basin because of a lack of a productive aquifer.

Surface-water-quality samples were collected from Symmes Creek (two sites), Indian Guyan Creek (two sites), Ice Creek, Little Indian Guyan Creek, Black Fork, Sand Fork, Johns Creek, Aaron Creek, Long Creek, and Swan Creek on October 1, 1985, between November 25 and 26, 1985, or on December 19, 1985 (fig. 12). Results of chemical analyses are presented in table 6 and figure 13.

Specific conductance ranged from 250  $\mu$ S/cm (at site T-3) to 465  $\mu$ S/cm (at site T-2), with a median of 370  $\mu$ S/cm.

Values of pH ranged from 6.70 (at site T-11) to 7.90 (at site T-7), with a median of 7.50. Alkalinity ranged from 10 mg/L as  $CaCO_3$  (at site T-11) to 182 mg/L (at site T-2), with a median of 94 mg/L.

Dissolved-sulfate concentrations ranged from 53 mg/L (at site T-1) to 170 mg/L (at site T-1), with a median of 81 mg/L.

Total-iron concentrations ranged from 350  $\mu$ g/L (at site T-5) to 1,000  $\mu$ g/L (at site T-6), with a median of 680  $\mu$ g/L. Dissolved-iron concentrations ranged from 20  $\mu$ g/L (at site T-9) to 150  $\mu$ g/L (at site T-1), with a median of 90  $\mu$ g/L. Concentrations of total aluminum ranged from 100  $\mu$ g/L (at sites T-2, T-5, and T-12) to 2,000  $\mu$ g/L (at site T-11), with a median of 200  $\mu$ g/L. Dissolved-aluminum concentrations were less than 100 or 100  $\mu$ g/L at all sites. Total-manganese concentrations ranged from 120  $\mu$ g/L (at site T-7) to 1,700  $\mu$ g/L (at site T-1), with a median of 235  $\mu$ g/L; dissolved-manganese concentrations ranged from 120  $\mu$ g/L (at site T-7) to 1,400  $\mu$ g/L (at site T-1), with a median of 230  $\mu$ g/L.

Date	Stream- flow, instan- taneous (ft <sup>3</sup> /s)	ci co du ta	n- c- nce	ph (stand- ard units)	Temper- ature (deg. C)	(mq/L	linity, field	sol ved	
38300508	2280600	T-1	Symme	s C nr (	etaway OH	(lat 38			
Oct 1985			410		15.0		105	53	200
38271508	2242400				nr Bradrio				82 24 24w)
									100
					llia OH (1				7w)
Nov 1985									
26	25				10.0			62	
		T-4	Symme	s C nr G	allia OH	lat 38 5	0 28n lon	g 082 29	
Nov 1985									
					9.5			994	300
384613082	2233600	T-5	Sand	F nr Pat	riot OH (1	at 38 46	13n long	082 23 3	6w)
Nov 1985					11.5				100
384125082	2283700	T-6	Johns	C nr Wa	terloo OH	(lat 38	41 25n lo	ng 082 28	
Nov 1985					10.5			67	200
					gus OH (la				w)
Nov 1985 26					10.5			71	200
					abia OH (1				 7w)
					11.0				
					ton OH (la				w)
Nov 1985 25	14				9.0			87	200
383332082	2205600	r-10	Indian	Guyan C	at Platfo	rm OH (1	at 38 33	32n long	082 20 56w
Dec 1985			435	7.30	0.5		87	140	200
383301087	2231400	r-11 1	L Indi	an Guyan	C at Scot	town OH	(lat 3833	Oln long	082 23 14w
Dec 1985	7.4				2.0		10	170	2,000
	124400			C nr Bl	aden OH (1	at 38 36	57n long	082 12 4	4w)
Nov 1985									

sites in the Symmes, Ice, and Indian Guyan Creeks basin

μg/L, micrograms per liter; μS/cm, microsiemens per centimeter at 25 degrees Celsiusl

Alum- inum-			Iron,		Manga-	Manga- nese,		
	lum-	Iron,	sus-		nese,	sus-	Manga-	
	num,		pended	Iron,	total	pended		
	dis-	recov-		dis-	recov-		dis-	
	olved		erable	solved			solved	
	µg/L		(µg/L	(µg/L	(µg/L	(ug/L	(µg/L	
	s A1)	as Fe)		as Fe)	as Mn)	as Mn)	as Mn)	
383005082280600	T-1	Symmes C	nr Getawa	ay OH (la	38 30 0	5n long 0	82 28 06w)	
	<100	860	710	150	1,700	300	1,400	
382715082242400		ndian Guya						24w)
	<100	360	220	140	550	100	450	
			0-114-	011 (1-4-1	00 50 26-	1 002	20 471	
385026082294700	T-3	Black F n	r Gallia	OH (lat .	38 50 26n	10ng 082	29 4/W)	
122	<100	850	780	70		20	460	
385028082293700			nr Gallia	OH (lat			2 29 37w)	
200	100	900	790	110	230	10	220	
384613082233600	T-5	Sand F nr		OH (lat		long 082	23 36w)	
0	100	350	300	50	220	0	220	
384125082283700	т-6	Johns C n	r Waterlo	o OH (lat	38 41 2	n long 0	82 28 37w)	
	<100	1,000	860	140	240	0	240	
383841082271400	T-7	Long C nr	Wilgus (	OH (lat 38	38 4ln	ong 082	27 14w)	
100	100	670	630	40	120	0	120	
383943082291700	т-8	Aaron C n	r Arabia	OH (lat 3	8 39 43n	long 082	29 17w)	
100	100	560	480	80	160	0	160	
3216050	T-9	Ice C at		OH (lat 38		long 082	38 29w)	
100	100	690	670	20	200	0	200	
383332082205600	T-10	Indian Guy	an C at I	latform (	H (lat 38	33 32n	long 082 2	0 56v
	<100	540	440	100	310	0	310	
383301082231400		L Indian G	uyan C at		OH (lat	38330ln	long 08223	14w)
	<100		830		1,300	a	1,300	
383657082124400	T-12			OH (lat		long 082	12 44w)	
집중하여 하는 사람이 아이를 하고 있다.								

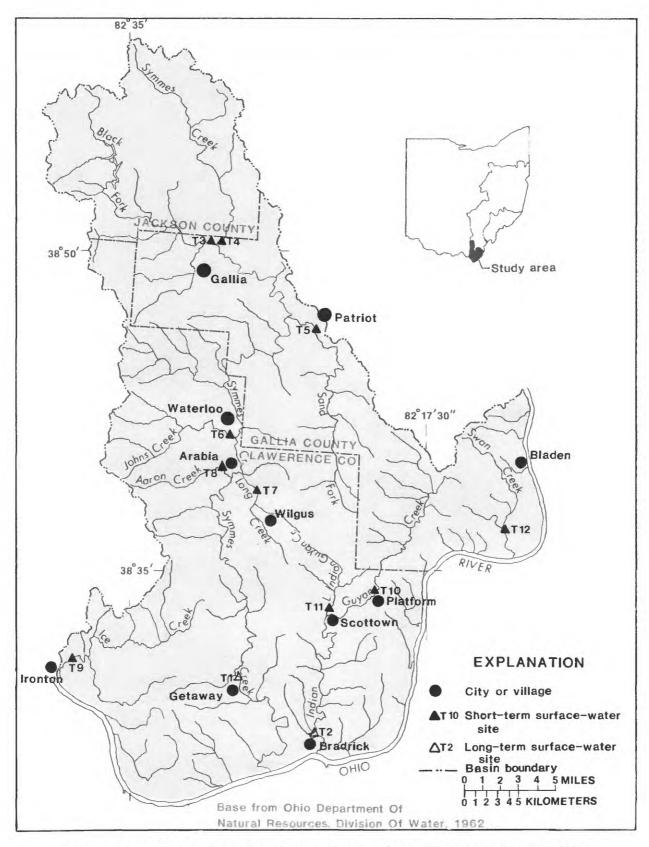


Figure 12.--Short-term and long-term surface-water sites in the Symmes, Ice, and Indian Guyan Creeks basin.

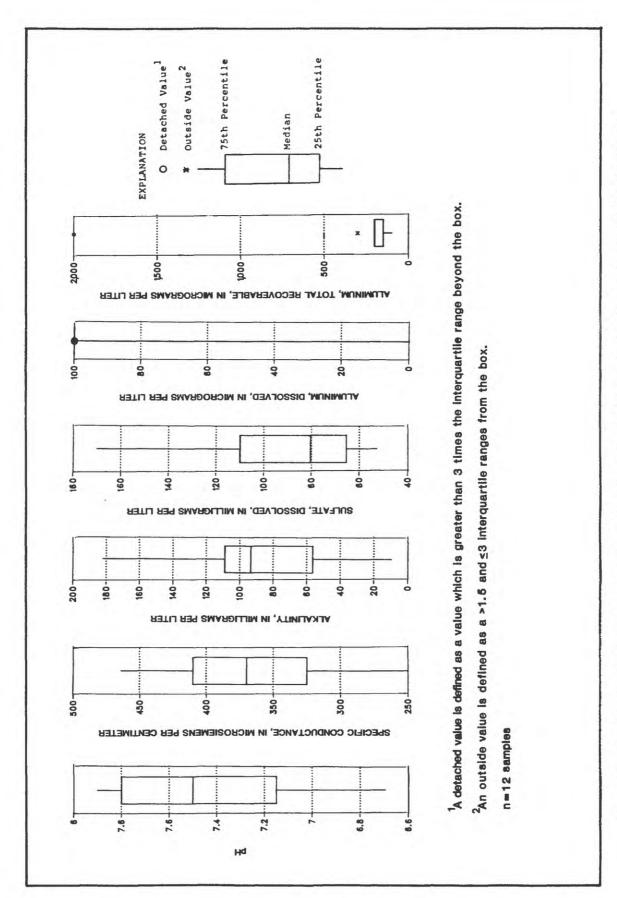
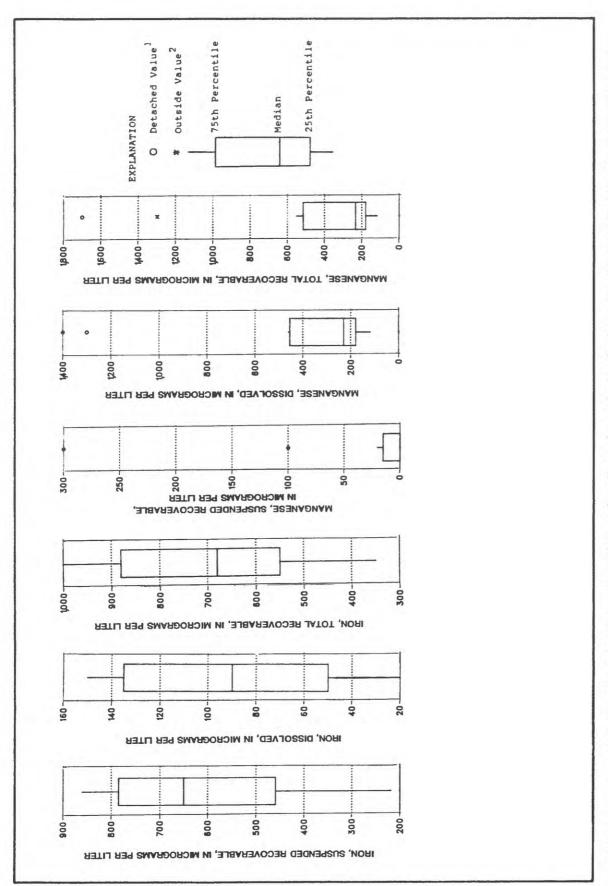


Figure 13.—Box plots showing the range, porcentiles, and median values of constituents at surface-water sites in the Symmes, ice, and indian Guyan Creeks basin.



constituents at surface-water sites in the Symmes, Ice, Figure 13.--Box plots showing the range, percentiles, and median values of and Indian Guyan Creeks basin--Continued,

#### Moxahala Creek Basin

### Geologic Setting

The geologic formations underlying the Moxahala Creek basin (fig. 2) consist of sandstone, shale, limestone, and coal of Mississippian and Pennsylvanian age and glacial deposits of clay, silt, sand, and gravel of Quaternary age.

Most of the basin is included in the Kanawha section of the Appalachian Plateaus physiographic province (Fenneman, 1938). There is a sharp contrast between the topography of the north-western part of the basin, in which glaciation makes the area one of subdued topography with a relief of about 100 feet, and the unglaciated southeastern part of the basin, where the topography is mature and the relief averages 200 feet.

The Moxahala Creek basin lies on the eastern flank of the Cincinnati Arch, which dips to the east-southeast at an average rate of 25 to 30 feet per mile. Therefore, the oldest rocks of Mississippian age crop out in the western and northwestern part of the basin, and the youngest beds, those of the Conemaugh Formation, are exposed in the eastern and southeastern part of the basin (Flint, 1951).

The northwestern part of the basin lies within the glaciated part of Ohio. The drift deposits average 30 feet in thickness over the bedrock in the upland areas, and are primarily clay containing lenses of sand and gravel. Most wells are drilled through the thin drift into the underlying sandstone and shale. Wells drilled in these bedrock formations generally yield adequate water supplies for farm and domestic needs. Although small industrial and municipal supplies have been reported from sandstones, yields of 5 to 10 gal/min are more typical of bedrock wells in the basin. In the unglaciated part of the basin, the depth to bedrock averages 10 feet. In the southeastern part of the basin, where shales predominate, supplies of less than 5 gal/min are to be expected. Many wells drilled in this area have inadequate yields, and owners must rely on shallow dug wells and cisterns (Walker, 1962b).

Deposits of sand, gravel, and clay that fill a buried valley beneath the present valley of Jonathan Creek are more than 200 feet thick in places. Wells in these deposits have yields of as much as 200 gal/min (Walker, 1962b). Additional drilling and aquifer tests would be needed to locate and fully define permeable horizons within this buried valley. Other unconsolidated deposits as much as 80 feet thick fill parts of the Moxahala and Jonathan Creeks valley (Walker, 1962b).

Ground water was investigated in the glacial outwash deposit that follows the valley of Jonathan Creek, a major tributary to Moxahala Creek (fig. 14). Seventeen water-level measurements were obtained and are listed in table 3. The available wells

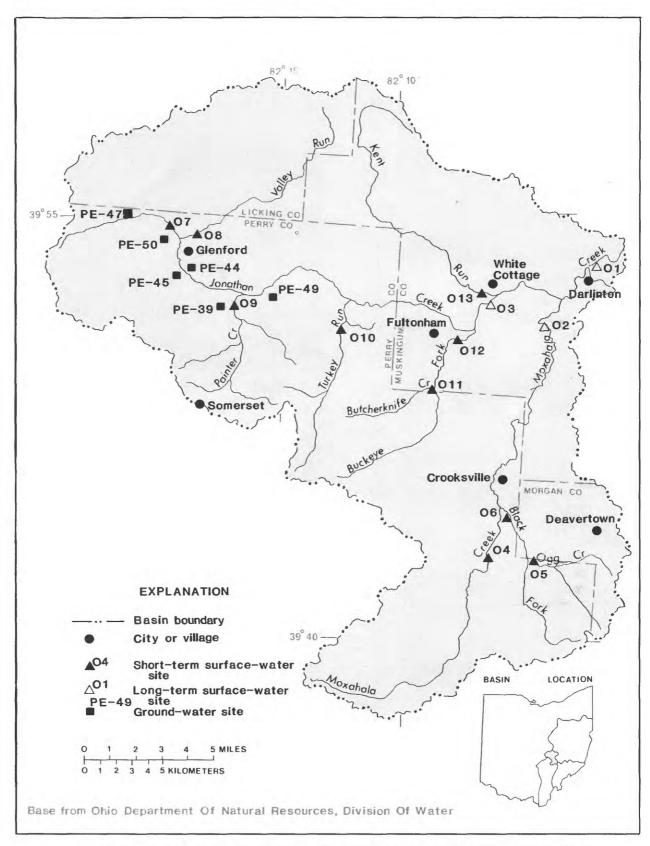


Figure 14.—Short-term and long-term surface-water sites and ground-water sites in Moxahala Creek basin.

completed in the sand and gravel did not extend far beyond the stream valley; therefore, it was impossible to draw a meaningful potentiometric surface with this limited number of observations, but it is likely that the local ground-water flow is toward the stream. Locations of wells are shown in figure 15.

## Water Quality

Surface-water quality of the Moxahala Creek basin was investigated by sampling 13 stream sites at low flow (fig. 14). The ground-water quality was investigated in the Jonathan Creek valley near Glenford, Ohio (fig. 15).

## Surface Water

Water-quality samples were collected from Moxahala Creek (three sites), Ogg Creek, Black Fork, Jonathan Creek (two sites), Valley Run, Painter Creek, Turkey Run, Butcherknife Creek, Buckeye Fork, and Kent Run during the period October 20 through October 22, 1986 (fig. 14). Results of chemical analyses are presented in table 7 and figure 16. Mining is presently active immediately adjacent to Ogg Creek in this basin.

Specific conductance ranged from 520  $\mu$ S/cm (at site O-9) to 4,600  $\mu$ S/cm (at site O-6), with a median of 1,600  $\mu$ S/cm.

pH ranged from 3.44 (at site O-2) to 9.20 (at site C-8), with a median of 6.60. Six samples had measurable acidity. The concentrations ranged from 9.0 mg/L as CaCO<sub>3</sub> (at site O-1) to 465 mg/L (at site O-5), with a median of 163 mg/L. Ten samples had measurable alkalinity. Alkalinity concentrations ranged from 2.0 mg/L as CaCO<sub>3</sub> (at site O-6) to 425 mg/L (at site O-8), with a median of 88 mg/L.

Dissolved-sulfate concentrations ranged from 42 mg/L (at site O-7) to 2,800 mg/L (at site O-5), with a median of 800 mg/L.

Total-iron concentrations ranged from 400 µg/L (at sites 0-7 and 0-9) to 110,000 µg/L (at site 0-5), with a median of 1,000 µg/L. Dissolved-iron concentrations ranged from 30 µg/L (at site 0-3) to 120,000 µg/L (at site 0-5), with a median of 500 µg/L. Total-manganese concentrations ranged from 90 µg/L (at site 0-9) to 25,000 µg/L (at site 0-4), with a median of 5,900 µg/L. Dissolved-manganese concentrations ranged from 80 µg/L (at site 0-9) to 23,000 µg/L (at site 0-4), with a median of 5,700 µg/L. Concentrations of total aluminum ranged from 70 µg/L (at site 0-9) to 26,000 µg/L (at site 0-5), with a median of 360 µg/L. Dissolved-aluminum concentrations ranged from 10 µg/L (at site 0-9) to 24,000 µg/L (at site 0-5), with a median of 80 µg/L. Water with a pH below 4.0 can contain several hundred or even several thousand milligrams of aluminum per liter.

Date	Stream- flow, instan- tangous (ft /s)	Spe- cific con- duc- tance (µS/cm)	ph (stand- ard units)	Temper- ature (deg. C)	Acidity (mg/L as CaCO <sub>3</sub> )	Alka- linity, field (mg/L as CaCO <sub>3</sub> )	Sulfate dis- solved (mg/L as SO <sub>4</sub> )	Alum- inum, total recov- erable (ug/L as Al)
39533708	2011100	0-1 Moxa	hala C nr	Darlingt	on OH (la	t 39 53 3	7n long	082 01 11w)
Oct 1986		1,600		12.0		25	800	360
03148400		0-2 Moxa	hala C at	Roberts	OH (lat 3	9 51 17n	long 082	03 23w)
Oct 1986		2,600					1,400	8,600
395214082	20554700	0-3 Jonati	han C at	White Cot	tage OH (	lat 39521	4n long (	
Oct 1986		1,050						
03148150	0-4	Moxahal	a C nr Cr	ooksville	OH (lat	39 43 52n	long 082	2 06 04w)
Oct 1986		2,600						
394340082	041200	0-5 Ogg (	nr Deav	ertown OH	(lat 39	43 40n lo	ng 082 04	1 12w)
Oct 1986		4,500						26,000
394519082	051600	O-6 Black	F nr Cr	ooksville	OH (lat	39 45 19n	long 082	05 16w)
Oct 1986		4,600						
395432082	194000	0-7 Jona	than C nr	Glenford	OH (lat	39 54 32n	long 082	19 40w)
Oct 1986 20		540				240		100
395419082	184400	O-8 Valle	y Rn nr (	Glenford	OH (lat 3	9 54 19n	long 082	
Oct 1986 20		540		8.0			60	
		O-9 Paint						16 56w)
Oct 1986		520					59	70
		0-10 Turk						
Oct 1986 20	0.39	1,000	7.65	8.5		56	390	110
394919082	082000	0-11 Butche				lat 39491	9n long 0	820820w)
Oct 1986 20	0.39	1,600		9.0			960	12,000
		0-12 Buckey	e F nr E	ast Fulto		lat 39504	8n long C	1820720w)
Oct 1986 22	1.8	1,600	6.60	12.0		13	900	1,200
		0-13 Ken	Rn at W	hite Cott	age OH (1	at 39 52		082 05 53w
Oct 1986 22		1,100					82	190

 $\mu g/L$ , micrograms per liter;  $\mu S/cm$ , microsiemens per centimeter at 25 degrees Celsiusl

			Iron,		Manga-	Manga-		
inum, sus-	Alum-	Iron,	SUS-		Manga-	nese,	Manga-	
pended	inum,	total		Tron	nese,			
recov-	dis-	recov-	pended	Iron, dis-	total	pended	nese,	
erable	solved	erable			recov-	recov- erable	solved	
(µg/L	(µq/L)		(µq/L	solved				
as Al)	as Al)	(µg/L) as Fe)	as Fe)	(µg/L) as Fe)		(µg/L) as Mn)	(µg/L) as Mn)	
395337082011					OH (lat 39			01 11w)
300	60	1,100	250	850	6,400	0	6,600	
03148400	0-2	Moxahala	C at Rob	erts OH	(lat 39 51	17n long	082 03	23w)
400	8,200	11,000	0	11,000	14,000	0	16,000	
395214082054	700 O-3 3	Jonathan C		Cottage		52 14n	long 082	05 47w
200	80	420	390	30	2,500	0	2,600	
03148150	0-4	Moxahala	C nr Cro	oksville	OH (lat 39	43 52n	long 082	06 04w
900	6,600	24,000	4,000	20,000	25,000	2,000	23,000	
394340082041	200 O-5	Ogg C nr	Deavertow	n OH (la	t 39 43 40n	long 08	2 04 12w	
2,000	24,000	11,000	0	12,000	9,300	0	9,500	
394519082051	600 O-6	Black F	nr Crooks	ville OH	(lat 39 45	19n lon	g 082 05	16w)
600	3,600				5,900			
395432082194	000 0-7	Jonathan	C nr Gle		(lat 39 54	32n lon	g 082 19	40w)
80	20	400	330	70	120	0	120	
			nr Glen	ford OH	(lat 39 54		082 18 4	4w)
395419082184	400 O-8	Valley R				19n long		
250	20	1,000	920	80	190	0	190	
250	20	1,000	920	80	190	0	190	6w)
250	20	1,000	920	80	190	0	190	 6w)
250 3952100821656	20 600 0-9	1,000 Painter (	920 C nr Some 320	80 rset OH	190 (lat 39 52	0 10n 10ng	190 082 16 5	
250 3952100821650 60 3951280821210	20 600 O-9 10 600 O-10	1,000 Painter 6 400 Turkey 1	920 C nr Some 320 Rn nr Som	80 rset OH 80 erset OH	190 (lat 39 52 90 (lat 39 51 2,300	0 10n long 10 28n lon	190 082 16 5 80 g 082 12 2,300	
250 8952100821650 60 8951280821210 90 8949190820820	20 600 O-9 10 600 O-10 20	1,000 Painter ( 400 Turkey ) 440 Butcherkn	920 C nr Some 320 Rn nr Som 380	80 rset OH 80 erset OH 60	190 (lat 39 52 90 (lat 39 51 2,300	0 10n long 10 28n lon	190 082 16 5 80 g 082 12 2,300	16w)
250 8952100821650 60 8951280821214 90 8949190820820	20 600 O-9 10 600 O-10 20	1,000 Painter ( 400 Turkey ) 440 Butcherkn	920 C nr Some 320 Rn nr Som 380	80 rset OH 80 erset OH 60	190 (lat 39 52 90 (lat 39 51 2,300	0 10n long 10 28n lon 0 94919n 1	190 082 16 5 80 g 082 12 2,300	16w)
90 3949190820820	20 600	1,000 Painter ( 400 Turkey ) 440 Butcherkn: 3,200 Buckeye F	920 C nr Some 320 Rn nr Som 380 ife C nr 400 nr East	80 rset OH 80 erset OH 60 Fultonham	190 (lat 39 52 90 (lat 39 51 2,300 n OH (lat 3 16,000	0 10n long 10 28n lon 0 94919n l	190 082 16 5 80 g 082 12 2,300 ong 08208	16w) 
250 3952100821650 60 3951280821210 90 3949190820820 1,000 3950480820720 570	20 600	1,000 Painter ( 400 Turkey ) 440 Butcherkn: 3,200 Buckeye F	920 C nr Some 320 Rn nr Som 380 ife C nr 400 nr East	80 rset OH 80 erset OH 60 Fultonham 2,800 Fultonham 500	190 (lat 39 52 90 (lat 39 51 2,300 DOH (lat 3 16,000 OH (lat 3	0 10n long 10 28n lon 0 94919n l 0 95048n l	190 082 16 5 80 g 082 12 2,300 ong 08208	16w) 
250 8952100821650 60 8951280821210 90 8949190820820 1,000 8950480820720 570	20 600 O-9 10 600 O-10 20 000 O-11 11,000 000 O-12 630	1,000 Painter 6 400 Turkey 1 440 Butcherkn: 3,200 Buckeye F	920 C nr Some 320 Rn nr Som 380 ife C nr 400 nr East	80 rset OH 80 erset OH 60 Fultonham 2,800 Fultonham	190 (lat 39 52 90 (lat 39 51 2,300 n OH (lat 3 16,000 n OH (lat 3	0 10n long 10 28n lon 0 94919n l 0 95048n l	190  082 16 5  80  g 082 12  2,300  ong 08208  18,000  ong 08207	16w) 

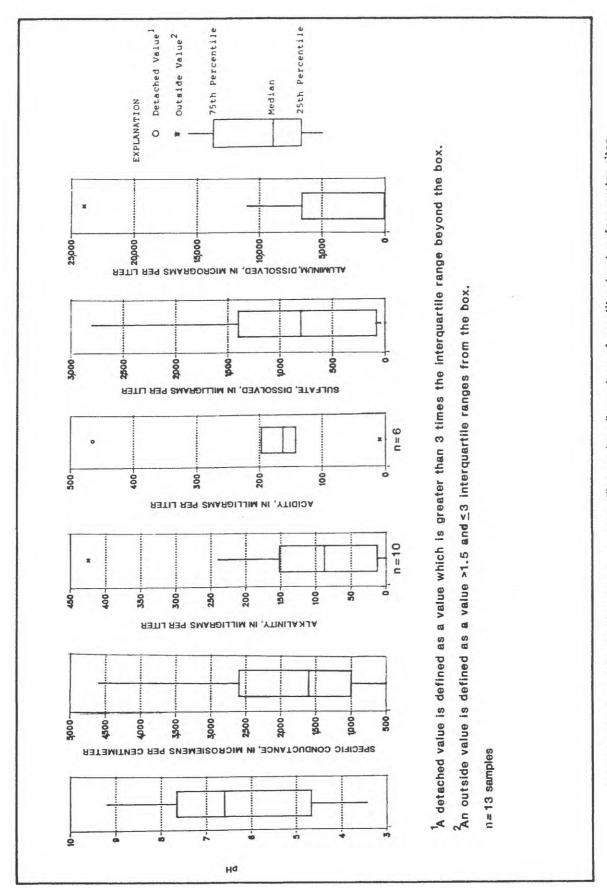


Figure 16.--Box plots showing the range, percentiles, and median values of constituents at surface-water sites in the Moxahala Creek basin.

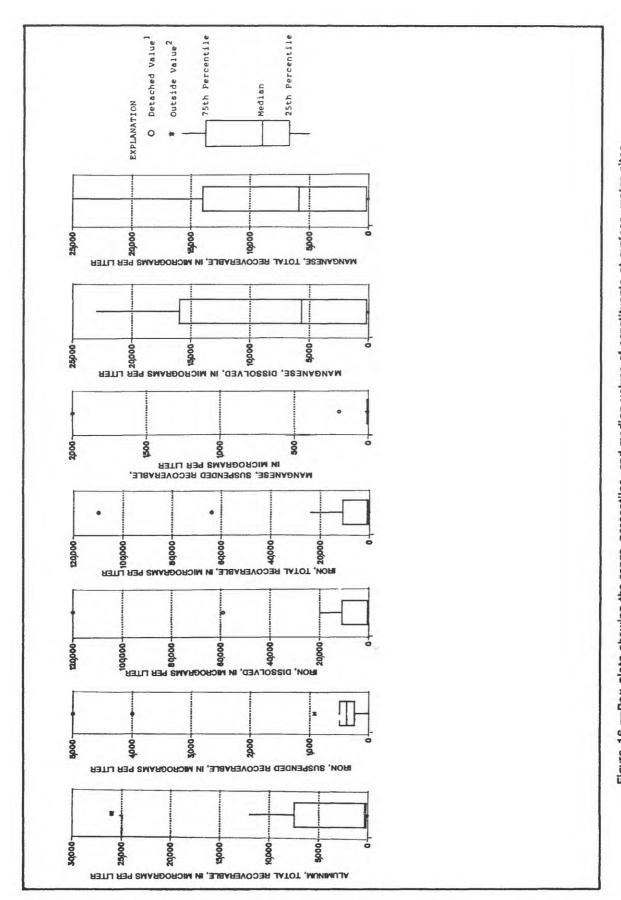


Figure 16.---Box plots showing the range, percentiles, and median values of constituents at surface-water sites in the Moxahala Creek basin--Continued.

Such water often occurs as drainage from mines (Hem, 1985). This is the case at sites O-1, O-2, O-4, and O-11 in the Moxahala Creek basin.

#### Ground Water

Ground-water samples were collected from five domestic wells completed in the sand and gravel aquifer at Glenford, Ohio, and vicinity in the Moxahala Creek basin (fig. 15). There are many active and abandoned mines in the Moxahala Creek basin, but, at present, only a few small mines are in operation southeast of Glenford. Table 8 lists the results of chemical analyses.

In general, the chemical quality is typical of other outwash aquifers in Ohio (Evans, 1977; deRoche and Razem, 1984; Norris and Fidler, 1969). This aquifer has a hard to very hard calcium bicarbonate type water with moderately high concentrations of dissolved solids. Specific conductance ranged from 447 to 670 µS/cm, and dissolved-solids concentrations ranged from 268 to 384 mg/L. Values for pH ranged from 7.20 to 7.80, and alkalinity ranged from 155 to 323 mg/L as CaCO<sub>3</sub>, both typical of outwash aquifers in Ohio.

Calcium is the most abundant cation and ranged from 52 to 87 mg/L in the five samples. Bicarbonate, the most abundant anion, ranged from 189 to 395 mg/L. Concentrations of magnesium ranged from 21 to 28 mg/L, and concentrations of sodium ranged from 12 to 23 mg/L. Hardness ranged from 220 to 330 mg/L as CaCO2, which classifies the water as hard to very hard. Potassium concentrations ranged from 1.1 to 4.5 mg/L. Concentrations of silica and chloride ranged from 10 to 15 mg/L and 5.2 to 40 mg/L, respectively. Total organic carbon in these wells, which are completed in sand and gravel, ranged from 0.6 to 1.0 mg/L, that are well within the average range for natural ground water. Three samples exceeded the OEPA standard for a public supply in dissolved-iron concentrations, and two were well within the limit (range, less than 3 to 4,400  $\mu g/L$ ). Total-iron concentrations ranged from 60 to 5,900 µg/L. Total-manganese concentrations ranged from less than 10 to 280  $\mu g/L$ , and three of the five samples exceeded the OEPA recommended criterion for a public supply. Dissolved-manganese concentrations ranged from less than 1 to 290 µg/L. Total-aluminum concentrations ranged from less than 10 to 890 µg/L; two samples greatly exceeded the concentration of the other samples. Dissolved-aluminum concentrations ranged from less than 10 to 20 µg/L. Clays are the most common of all the sedimentary aluminum-bearing minerals. The polymerization of aluminum hydroxide species proceeds in a different way in the presence of silica. When sufficient silica is present, the aluminum appears to be rapidly precipitated (Hem, 1985). Lenses of clay appear throughout this outwash aquifer. In the Moxahala Creek basin, the samples from wells for which logs indicate considerable amounts of clay generally had aluminum concentrations that were above average and silica concentrations that were slightly below average.

[deg. C, degrees Celsius; mg/L, milligrams per liter;  $\mu$ g/L, micrograms per liter;  $\mu$ S/cm, microsiemens per centimeter at 25 degrees Celsius]

Date	level)	duc- tance	ph (stand- ard units)	Temper- ature (deg. C)	sol ved	(mg/L as	bonate, field	solved (mg/L	dis- solved (mg/L
	2190600 PE-								
	36.87								
395309082	2192500 PE-	-45 (lat	39 53 09n	long 082	2 19 25w)				
08/19/86		480	7.20	13.0		220	65	52	22
395507082	2210500 PE-	-47 (lat	39 55 07n	long 082	21 0561				
	0.0								
395218082	2151700 PE-	-49 (lat	39 52 18n	long 082	15 17w)				
08/22/86	1.08	573	7.50	14.0	0	250	0	62	23
395431082	195500 PE-	-50 (lat	39 54 31n	long 082	19 55w)				
					3.1	330		87	28
Date	Sodium, dis- solved (mg/l as Na)	Potas- sium, dis- solved (mg/L as K)	bonate,	linity, field	dis- solved	dis- solved	dis- solved (mg/L	deg. C	
395320082	190600 PE-	44 (lat	39 53 20n	long 082	19 06w)				
08/18/86	13	4.5	329	270	47	13	14	32	9
395309082	192500 PE-	-45 (lat	39 53 09n	long 082	19 25w)				
08/19/86	12	1.5	189	155	34	40	10	27	1
	210500 PE-								
	17								
	151700 PE-	49 (lat	39 52 18n	long 082	15 1				
08/22/86	23	3.0	312	256	63	5.2	13	271	8
	195500 PE-								
08/22/86	18	1.9	343	281	36	36	15	38	4
Date	Solids, sum of constit- uents, dis-	Alumi- num, total recov- erable (ug/L	Alumi- num,	Alumi- num, sus- pended recov- erable (ug/L	Iron, total recov- erable (ug/L	Iron,	Manga- nese, total	Manga- nese, dis-	Carbon, organic total (mg/L as C)
395320082	190600 PE-	44 (lat	39 53 20n	long 082	19 06w)				
08/18/86	360	40	<10		60	6	<10	<1	1.0
395309082	192500 PE-	45 (lat	39 53 09n	long 082	19 25w)				
08/19/86	270	20	<10		690	620	60	57	0.6
395507082	210500 PE-	47 (lat	39 55 07n	long 082	21 05w)				
08/18/86	320	890	<10	1	,500 ]	,400	280	290	1.0
205218082	151700 PE-	49 /1a+	39 52 18n	long OR2	15 17w1				
8/22/86	350	480	20	460 5	,900 4	,400	170	290	0.7
205421082	105500 DE-	50 (1at	30 54 31n	long 082	19 5561				
0.422406	300	<10	<10		80	<3	<10	1	0.6

#### Little Beaver Creek Basin

## Geologic Setting

The Little Beaver Creek basin (fig. 2) is located on the stream-dissected Appalachian Plateau (Fenneman, 1938). The bedrock consists of sandstone, siltstone, clay, shale, and limestone of Pennsylvanian age. These rocks were elevated several times, thus, the present land surface is a series of stepped levels, each a remnant of sporadic erosion. The high valleys have been entrenched in the eastern part of the basin through changes in drainage brought about by Pleistocene glaciation in the northern part of the basin. These processes formed the narrow valleys and gorges in the basin that range in elevation from 920 feet at the headwaters of Little Beaver Creek to 650 feet along the Ohio River (Lessig and others, 1968).

The sedimentary rocks underlying the Little Beaver Creek basin are relatively impermeable sandstone and shale of the Allegheny and Conemaugh Formations. The southern part of the basin is hilly and rugged. Here, the Allegheny rocks crop out in the valley walls and dip about 8 feet per mile to the southeast. On uplands in the southern part of the basin, most of the rocks exposed belong to the lower part of the Conemaugh Formation. This part of the basin, the unglaciated section, has a potential ground-water yield of less than 5 gal/min (Lessig and others, 1968).

Little Beaver Creek has a limited area for the development of large ground-water supplies. Wells completed in the glacial deposits beneath the narrow flood plain of the Ohio River have an estimated potential yield of 500 to 1,000 gal/min. There are little accurate data that reveal this potential; however, some driller's logs report as much as 42 feet of sand and gravel beneath a thin layer of alluvium (Lessig and others, 1968).

Sand and gravel are abundant in the northern part of the basin. Many quarries have operated in the past, and a number of smaller pits are operated intermittently. Although most of the drilled wells are developed in the bedrock, many well logs reveal lenses of sand and gravel above bedrock. Wells in these unconsolidated deposits have a potential yield of 5 to 25 gal/min. Where these deposits are thin, the underlying bedrock may yield up to 25 gal/min (Schmidt, 1959a).

Uplands in the northern part of the basin are covered by thick glacial till of Wisconsin age ranging from 10 to more than 160 feet in thickness. Till and outwash of Wisconsin age are deposited along Little Beaver Creek and its tributaries. About 70 percent of this material is of local origin, mostly sedimentary rocks. The rest has been transported from Canada and other areas to the north. Clay is more abundant in the northwestern part of the basin (Lessig and others, 1968).

The glacial fill in the discontinuous valleys beneath some of the minor streams consists of thick deposits of clay interbedded with thin to thick layers of fine sand and gravel. Yields of as much as 65 gal/min have been reported, yet favorably sited wells having yields of as much as 100 gal/min appear possible (Schmidt, 1959a). Sand and gravel deposits with the greatest potential yield are the preglacial and interglacial drainage channels, which are now wholly or partially filled with glaciofluvial deposits of varying thickness. The complex valleys of Middle Fork Little Beaver Creek, and East Branch Middle Fork Little Beaver Creek, in Columbiana County, also have large deposits that contain water (White and Totten, 1985).

Ground-water levels were measured in 23 wells in the area near Lisbon (fig. 17), and are presented in table 3. Locations of wells are shown in figure 18.

## Water Quality

Surface-water quality of the Little Beaver Creek basin was investigated by sampling 12 stream sites at low flow (fig. 17). Ground-water quality was investigated by sampling five wells near Lisbon (fig. 18).

#### Surface Water

Water-quality samples were collected from Little Beaver Creek, Middle Fork Little Beaver Creek (two sites), North Fork Little Beaver Creek, West Fork Little Beaver Creek, East Branch Middle Fork Little Beaver Creek, Bull Creek, Longs Run, Cold Run, Brush Creek, Elk Run, and Little Yellow Creek during the period October 21 through October 23, 1986 (fig. 17). Results of chemical analyses are presented in table 9 and figure 19.

Specific conductance ranged from 330  $\mu$ S/cm (at site A-6) to 1,210  $\mu$ S/cm (at site A-3), with a median of 665  $\mu$ S/cm.

Values for pH ranged from 7.49 (at site A-12) to 8.83 (at site A-1), with a median of 8.37. Alkalinity ranged from 35 mg/L as  $CaCO_3$  (at site A-12) to 257 mg/L (at site A-4), with a median of 153 mg/L.

Dissolved-sulfate concentrations ranged from 41 mg/L (at site A-6) to 250 mg/L (at site A-11), with a median of 145 mg/L.

Total-iron concentrations ranged from 150  $\mu$ g/L (at site A-5) to 2,700  $\mu$ g/L (at site A-3), with a median of 410  $\mu$ g/L. Dissolved-iron concentrations ranged from 20  $\mu$ g/L (at site A-12) to 80  $\mu$ g/L (at site A-9), with a median of 50  $\mu$ g/L.

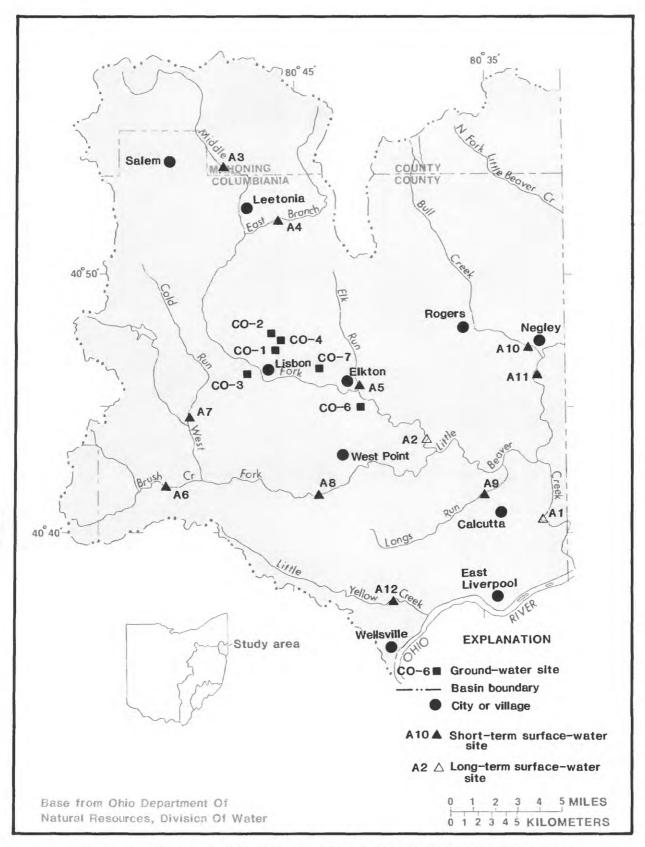


Figure 17.—Short-term and long-term surface-water sites and ground-water sites in the Little Beaver Creek basin.

Total-manganese concentrations ranged from 50  $\mu$ g/L (at site A-1) to 780  $\mu$ g/L (at site A-3), with a median of 85  $\mu$ g/L. Dissolved-manganese concentrations ranged from 20  $\mu$ g/L (at site A-12) to 760  $\mu$ g/L (at site A-3), with a median of 55  $\mu$ g/L.

Concentrations of total aluminum ranged from 20  $\mu g/L$  (at site A-10) to 390  $\mu g/L$  (at site A-7), with a median of 125  $\mu g/L$ . Dissolved-aluminum concentrations ranged from less than 10  $\mu g/L$  (at sites A-9 and A-10) to 30  $\mu g/L$  (at site A-12), with a median of 20  $\mu g/L$ .

#### Ground Water

Ground water in the Little Beaver Creek basin was investigated in the vicinity of Lisbon, Ohio (fig. 17), mainly in the shallow sandstone of the Conemaugh Formation. One sample was collected from an outwash glacial deposit north of the village (well CO-2) for comparison with the bedrock water. The northwestern and central parts of the basin are covered with glacial deposits, but most are relatively thin; water users generally drill through the glacial cover to the underlying sandstone. The Village of Lisbon has a public water supply north of town, and one of these wells (CO-4) was used for sampling. The other four samples were collected from domestic wells in the Table 10 lists the results of chemical analyses. Surface mining is present in much of the basin, although, in general, it is characterized by groups of small mines rather than by sweeping coverage of the entire area. Well locations and Stiff waterquality diagrams are shown in figure 18.

A very hard calcium bicarbonate type water is present in the Lisbon area. In general, the water quality in this area is typical of bedrock aquifers in eastern Ohio (Razem and Sedam, 1985). Water quality of the sample from the outwash aquifer (CO-2) was similar to the bedrock water quality but was slightly higher in iron and manganese. One sample (CO-1) had higher concentrations of the metallic ions and dissolved solids and lower pH. This well, located downgradient from a nearby surface-mining operation, could possibly have received contaminants through the aquifer. The water from well CO-1 was classified as a calcium sulfate type.

Overall, specific conductance ranged from 51l to 1,300  $\mu$ S/cm. Dissolved-solids concentrations ranged from 293 to 1,050 mg/L. Only well CO-l exceeded the OEPA standard for a public supply in these constituents. Values for pH ranged from 6.65 to 7.20. Alkalinity ranged from 210 to 365 mg/L as CaCO<sub>3</sub>.

(deg. C, degrees Celsius; ft3/s, cubic feet per second; mg/L, milligrams per liter;

Date	Stream- flow, instan- taneous (ft)	tano	ic - - ce		Temper- ature	(mg/L as	linity, field (mg/L as	Sulfate dis- solved (mg/L as SO <sub>4</sub> )	recov- erable (µg/L
03109500	A-1	L Bear	ver C	nr East	Liverpool	OH (lat	40 40 3	n long 08	30 32 27w)
Oct 1986 22					12.5		130	160	90
03109100		A-2 M	FL	Beaver C	nr Rogers	OH (lat	40 43 2	2n long Of	30 38 03w)
Oct 1986 21	27		765	8.81	12.0		147	150	130
		A-3 1	MFL	Beaver	C nr Salem	OH (lat	40 54 20	n long 08	30 48 17w)
Oct 1986 21	7.7	1,	210	7.73	8.0		158	160	219
03108990	A-4 I	EBMI	FLB	eaver C	at Leetoni	a OH (la	t 40 52 1	len long (	80 45 54w)
Oct 1986 21	2.0		790	8.14	7.5		257	140	200
404544080	415400	A-5 1	Elk R	n at Elk	ton OH (la	t 40 45	44n long	080 41 54	
Oct 1986 21			550	8.76			138	170	30
404204080	515600	A-6 E	Brush	C nr We	st Point O	H (lat 4	0 42 04n	long 080	
Oct 1986 21					11.0			41	120
404423080	502900	A-7 (	Cold	Rn nr Li	sbon OH (1	at 40 44			9w)
Oct 1986 21				8.44	10.0		158	79	390
03109200	A-8	WFI	Bea	ver C at	West Poin	t OH (la	t 40 42 3	8n long 0	80 41 49w)
Oct 1986 22					9.0				
404140080	351100	A-9 I	Longs	Rn nr C	alcutta OH	(lat 40	41 40n 1	ong 080 3	5 11w)
Oct 1986					10.0			109	
03109395		A-10	Bull	C at Ne	gley OH (1	at 40 47	15n long	080 32 4	2w)
Oct 1986 22	11		570	8.30	9.0		170	109	20
03109400		A-11	NF	L Beaver	C nr Negl				080 32 36w
Oct 1986 22	29		930	8.44			169	250	50
403715080	391400	A-12	L Ye	llow C n	r Wellsvil	le OH (1	at 40 37	15n long	080 39 14w
Oct 1986								230	180

## surface-water sites in the Little Beaver Creek basin

 $\mu g/L$ , micrograms per liter;  $\mu S/cm$ , microslemens per centimeter at 25 degrees Celsiusl

Alum- inum,			Iron,		Manga-	Manga- nese,		
sus-	Alum-	Iron,	sus-		nese,	sus-	Manga-	
pended	inum,	total	pended	Iron,	total	pended	nese,	
recov-	dis-	recov-	recov-	dis-	recov-		dis-	
erable	solved	erable	erable	solved	erable	erable	solved	
(µg/L	(µq/L	(µq/L	(µg/L	(µg/L	(µg/L	(µg/L	(µg/L	
as Al)	as Al)	as Fe)	as Fe)	as Fe)	as Mn)	as Mn)	as Mn)	
3109500	A-1 B	eaver C n	r East Li	verpool OH	(lat 40	40 33n 1	ong 080 32	27w)
80	10	340	290	50	50	20	30	
3109100	A-2			Rogers OH			ong 080 38	03w)
110	20	560	530	30	80	40	40	
03108980							ng 080 48	17w)
199	20	2,700	2,700	50	780	20	760	
03108990 A-	4 E B			Leetonia O		0 52 16n	long 080 4	5 54w
190	10	1,100	1,100	40	300	90	210	
40454408041540				OH (lat 40				
10	20	150					109	
40420408051560		Brush Cr	nr West	Point OH (	lat 40 4			Бw)
100	20	720	660	60	90	40	50	
40442308050290	0 A-7	Cold Rn	nr Lisbon	OH (lat 4	0 44 23n	long 080	50 29w)	
380	10	1,300	1,200	50	160	50	109	
03109200	A-8 W	F L Beave	r C at We	st Point O	H (lat 4	0 42 38n	long 080 4	1 49
120	20	480	450	30	90	30	60	
10414008035110	0 A-9	Longs Rn	nr Calcu	tta OH (la	t 40 41	40n long	080 35 11w	)
	<10	320	240	80	80	40	40	
33109395	A-10	Bull C	at Negley	OH (lat 4	0 47 15n	long 080	32 42w)	
	<10	190	140	50	70	0	70	
03109400	A-11	N F L Be	aver C nr	Negley OH	(lat 40	46 30n l	ong 080 32	36w)
30	20	290	250	40	60	10	50	
40371508039140	0 A-12	L Yellow	C nr Wel	lsville OH	(lat 40	37 15n 1	ong 080 39	14w)

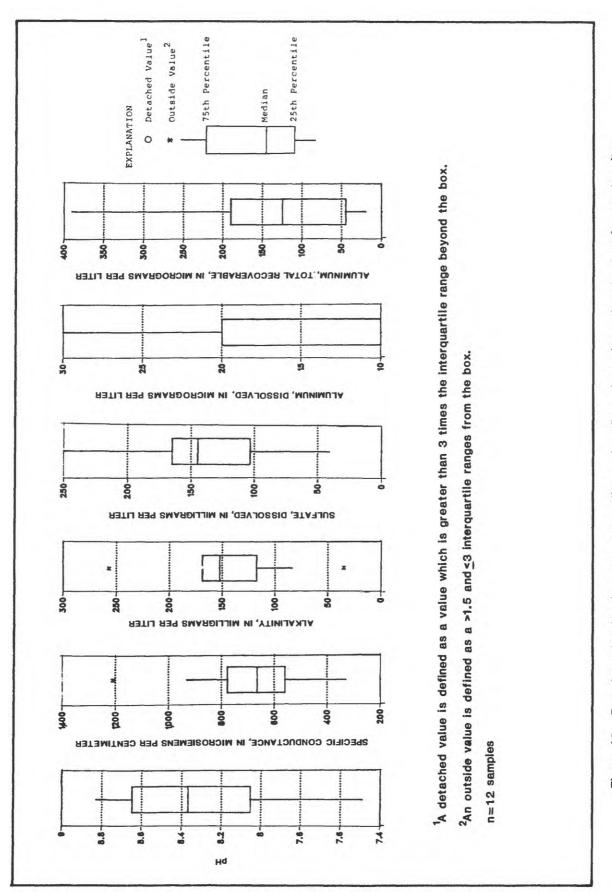


Figure 19.--Box plots showing the range, percentiles, and median values of constituents at surface-water sites in the Little Beaver Creek basin,

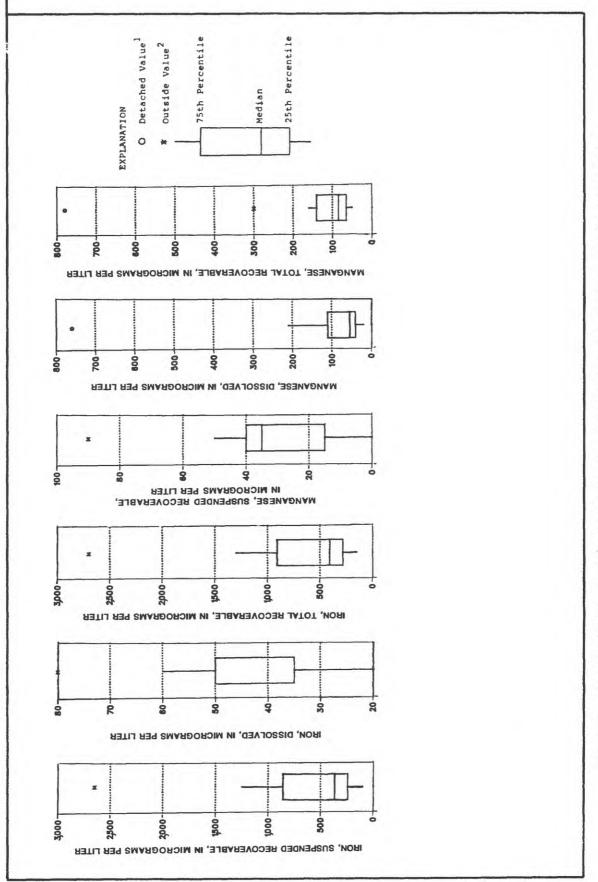


Figure 19.-Box plots showing the range, percentiles, and median values of constituents at surface-water sites in the Little Beaver Creek basin--Continued.

# Table 10.--Ground-water analyses for the Little Beaver Creek basin, Lisbon, Ohio, and vicinity

[deg. C, degrees Celsius; mg/L, milligrams per liter; ug/L, micrograms per liter; uS/cm, microsiemens per centimeter at 25 degrees Celsius]

Date	Depth below land surface (water level) (feet)	duc-	(stand- ard	Temper- ature (deg. C)	solved	ness (mg/L as	bonate, field (mg/L as	Calcium dis- solved (mg/L as Ca)	(mg/L
404754080	453700 CO-	l (lat	40 47 54n	long 080	45 37w)				
	44.87								64
	460800 CO-2	(lat	40 50 08n	long 080	46 08w)				
09/03/86	0.0	750	7.10	11.5	0.1	440	65	120	33
404639080	483300 CO-	11at	40 46 39n	100g 080	48 33wl				
09/02/86	44.20	570	6.80	14.0	0	320	99	90	22
101916090	461000 00-	11126	AD AQ A6-	1000 000	46 7000				
09/03/86	23.00	512	7.20	10.5	0.3	260	0	75	18
104506080	414100 CO-6	(lat	40 45 06n	long 080	41 41w)				
	47.87						110	98	27
Date	Sodium, dis- solved (mg/L as Na)	Potas- sium, dis- sol ved (mg/L as K)	Bicar- bonate, field (mg/L as HCO <sub>3</sub> )	Alka- linity, field (mg/L as CaCO <sub>3</sub> )	Sulfate dis- solved (mg/L as SO <sub>4</sub>	Chlo- ride, dis- solved (mg/L ) as Cl)	Silica, dis- solved (mg/L as Sio <sub>2</sub> )	Solids, residue at 180 deg. C dis- solved (mg/L)	
	453700 CO-1								
						10	15	1050	
05008080	6.5 460800 CO-2	(lat	40 50 08n	long 080	46 08w)				
						4.2	19	304	
04639080	483300 CO-3	(lat	40 46 39n	long 080	48 33w)				
9/02/86	4.6	1.3	260	214	100	8.2	8.7	315	
04846080	461000 CO-4	(lat	40 48 46n	long 080	46 10w)				
	13								
	43 43 00 00 /	11-1	AO AE OC-	1 000	41 47)				
09/03/86						2.2	15	408	
Date	Solids, Assum of reconstittion uents, dissolved (mg/L) as	num, cotal cecov- erable (ug/L	Alumi- num, 1 dis- solved (ug/L as Al)	(ug/L (u as Al) as	ig/L (	ron, to dis- re olved en ug/L (to s Fe) as	cov- cable so	dis- orgolved to ug/L (m	bon, anic tal g/L C)
404754080	453700 CO-1	(lat	40 47 54n	long 080	45 37w)				
09/02/86	1,000	20	<10	2,6	500 2,	500 3,	200 3,	200 0	. 8
05008080	460800 CO-	2 (lat	40 50 08n	long 080	46 08w)	222 - 22			
09/03/86	490	40	<10	1,7	700 1,	700	390	390 0	. 4
04639080	483300 CC-	(lat	40 46 39n	long 080	48 33w)				
09/02/86	360	30	<10	9	80	140	60	57 6	.0
A D A D A C DOD	ACTOOD OO	1100	AO AO ACA	1000 000	46 1 Dust				
09/03/86	350	30	<10	{	390	800	250	240 0	. 4
ADAEDEDOD	AT 41 00 CO-	6 /1 24	40 45 DEn	long 080	41 41w)				
	420	80	<10	5	60	13	90	57 0	.5

Calcium is the most abundant cation; concentrations ranged from 75 to 210 mg/L for the five samples. Bicarbonate is the most abundant anion; concentrations ranged from 260 to 450 mg/L. Concentrations of magnesium ranged from 18 to 64 mg/L, and concentrations of sodium ranged from 4.6 to 13 mg/L. Hardness ranged from 260 to 790 mg/L as CaCO<sub>3</sub>, which classifies the water as hard to very hard. Potassium concentrations ranged from 1.3 to 2.2 mg/L. Concentrations of silica and chloride ranged from 8.7 to 19 mg/L and 2.2 to 10 mg/L, respectively. Total organic carbon ranged from 0.4 to 0.8 mg/L in four wells but was elevated to 6.0 mg/L in well CO-3.

Concentrations of total iron ranged from 560 to 2,600  $\mu g/L$ , and concentrations of dissolved iron ranged from 13 to 2,500  $\mu g/L$ . Three of the five samples exceeded the OEPA standard for a public supply. Total-manganese concentrations (60 to 3,200  $\mu g/L$ ) exceeded the OEPA limit for a public supply in all 5 wells. Total-aluminum concentrations ranged from 20 to 80  $\mu g/L$ , whereas dissolved-aluminum concentrations were less than 10  $\mu g/L$  in all samples.

## McMahon, Captina, and Sunfish Creeks Basin

## Geologic Setting

The McMahon, Captina, and Sunfish Creeks basin (fig. 2) is in the unglaciated, dissected Appalachian Plateau region of Ohio (Fenneman, 1938). The eastern part of the basin is the most rugged. Ridgetops and valleys are narrow, and side slopes are steep. To the west and along drainage divides the basin is not quite as rugged. In this area, ridgetops are wider and more gently sloping, and hills are less steep. The western part of the basin has been extensively surface mined. The average relief in the area is 700 feet. All three streams in this basin flow into the Ohio River. The rock strata in the basin are comprised of the Conemaugh and Monongahela Formations of Pennsylvanian age and the Dunkard Group of Pennsylvanian and Permian age. These layers generally dip to the southeast with an average slope of 18 feet per mile (Hayhurst and others, 1974).

The Conemaugh Formation, principally limestone and sandstone, is the least exposed formation in the basin. It crops out only in the extreme western part of the basin and in the valleys on the northeastern side of the basin. The Monongahela Formation, which consists of shale, limestone, sandstone, and coal, is located in the western part of the basin and along the major streams. The Dunkard Group is extensively exposed in the basin and crops out predominantly in the eastern and central parts of the basin. It consists of a variable sequence of limestones, sandstones, and shales (Rubel and others, 1981).

Less than 5 percent of the basin is suitable for the development of large industrial or municipal ground-water supplies. The glacial outwash deposits beneath the narrow flood plain adjacent to the Ohio River may yield 500 to 1,000 gallons per minute to properly developed wells. Accurate data to substantiate this potential is scarce, but the logs of drilled wells in the basin are similar to those of wells having similar yields developed in the permeable outwash deposits found elsewhere in the Ohio River flood plain (Schmidt, 1959b).

The potential yield of wells drilled beyond the influence of the Ohio River is 25 to 100 gal/min. Although these wells penetrate unconsolidated deposits similar to those beneath the Ohio River, yields are limited by the location with respect to recharge and proximity to the bedrock valley wall (Schmidt, 1959b).

More than 90 percent of the basin is underlain by bedrock formations that provide relatively low yields to wells. Yields of less than 2 gal/min are developed in the various layers of sandstone, shale, and limestone. The variation in the physical characteristics of the bedrock and the topography influences the yields of wells developed in this area. The most productive bedrock wells are developed in the valleys adjacent to the streams. Locally, sandstones may yield as much as 5 gal/min or more. However, the rapid runoff of precipitation from the rugged hill-sides generally directs the infiltration of recharge to the more permeable deposits along valley drainage (Schmidt, 1959b).

The Holocene alluvium deposits on the broader flood plains of McMahon, Captina, and Sunfish Creeks often consist of fine sand, gravel, silt, and clay. Where these materials are permeable, they may yield as much as 25 gal/min to large-diameter drilled or dug wells. In addition, these alluvial deposits, even though relatively thin, can provide recharge to the underlying bedrock and thereby influence yields of wells developed in that material (Schmidt, 1959b).

#### Water Quality

Surface-water quality of the McMahon, Captina, and Sunfish Creeks basin was investigated by sampling 13 stream sites at low flow (fig. 20). The ground-water quality was not investigated because a productive aquifer is absent.

Surface-water-quality samples were collected from McMahon Creek (two sites), Captina Creek, Sunfish Creek (two sites), South Fork Captina Creek, North Fork Captina Creek, Williams Creek, Little McMahon Creek, Bend Fork, Pea Vine Creek, Baker Fork, and Piney Fork, from October 20 through October 22, 1986 (fig. 20). Results of chemical analyses are presented in table 11 and figure 21.

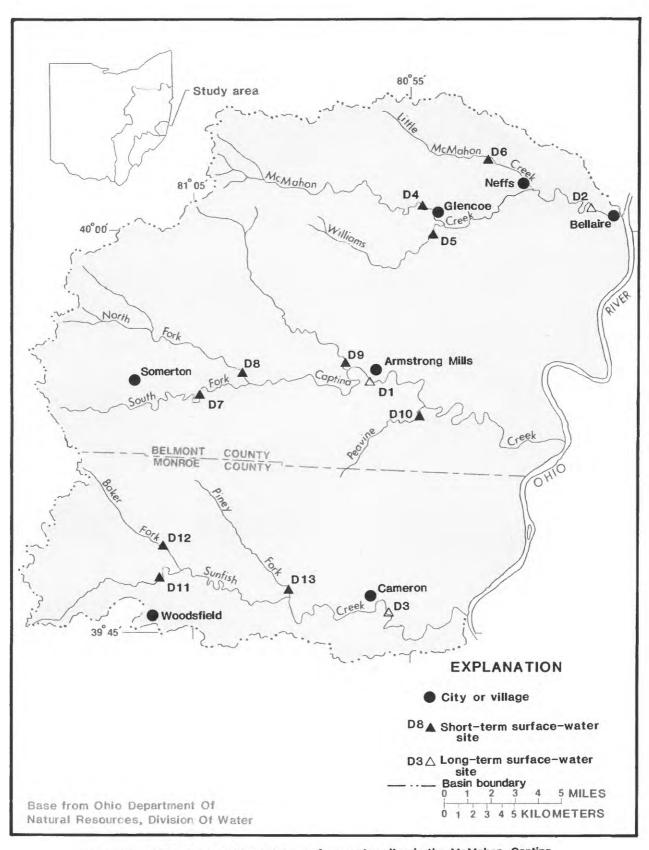


Figure 20.--Short-term and long-term surface-water sites in the McMahon, Captina, and Sunfish Creeks basin.

[deq. C, degrees Celsius; ft<sup>3</sup>/s, cubic feet per second; mg/L, milligrams per liter;

Date	flow, instan- taneous (ft <sup>3</sup> /s)	ci co du ta	c- nce	ph (stan- dard units)	Temper- ature (deg. C)	Acidity (mg/L as CaCO <sub>3</sub> )	linit	d as	Sulfated dis- solve (mg/las SO)	ed L	rec era (µg	ov-
03114000					rmstrongs l	dills OH	(lat 3	9543	ln lo		05	27w)
Oct 1986												
21	13		605	8.39				41	120			120
03113550		D-2		on C at	Bellaire (	OH (lat	40 00 3					
Oct 1986 20			980	8.31				46	330			210
					Cameron Ol	(lat 3	9 46 00	n lo	ng 080	56	094	)
Oct 1986												
22	4.3		435	8.31	10.5		1	25	52			70
40002308	532000	D-4			Glencoe O	(lat 4				53	20	)
Oct 1986 21	9.6			8.63				34	160			90
		D-5	Willi	ams C a	t Glenco Ol	(lat 4	00 13	n lo	ng 080		30w	)
Oct 1986 21	1.2		485		13.5			53				50
400225080	504100	D-6	L McM	ahon C	nr Neffs OF	(lat 4)	0 02 25	n lo	ng 080	50	41w	)
Oct 1986												
20	1.8	1,	,200	7.94	12.0		1	92	410		3,	500
395419081	044800	D-7 S	S F Car	otina C	nr Somerto	n OH (1a	at 39 5	4 19	n long	081	04	4 8w
Oct 1986 21	2.0				9.5		1					170
395444081	025000	D-8	N F C	aptina C	Somerton	OH (lat	39 54	44n	long 0	81 0	2 5	0w)
Oct 1986 21	4.5		550	8.75	8.5		1	65	79			110
955020805	75700 D	-9 Ber	nd Fk	nr Armst	trongs Mill	s OH (1a					57	57w)
Oct 1986 21	2.4		455					40				30
	41300 D	-10 Pe	a Vin	e C nr A	Armstrongs	Mills OF						
Oct 1986 21	0.41			8.33	7.5		1	86	59			40
394712081			Sunf		Woodsfiel						07	01w)
Oct 1986 22	0.28			8.78	10.0			05	51			40
					Woodsfield							53w)
Oct 1986 22	0.27		350	8.12	9.0			85	38			50
394645081	004100				Woodsfield					081	00	41w)

sites in the McMahon. Capting, and Sunfish Creeks basin

	rams per	liter; p	s/cm, mici	ostemens	per centi	meter at	25 degreea ce	tatuai
Alumi-						Manga-		
num,	33.v=1		Iron,		Manga-	nese,	Manga-	
sus-	Alumi-	Iron, total	sus- pended	Iron,	nese, total	nended	Manga- nese,	
pended recov-	num, dis-	recov-		dis-	recov-		dis-	
erable	solved			solved	erable	erable	solved	
(uq/L	(µg/L		(µg/L	(µg/L	(ug/L	(µg/L	(ug/L	
as Al)	as Al)	as Fe)	as Fe)	as Fe)	as Mn)	as Mn)	as Mn)	
3114000	D-1	Captina (	C at Armst	rongs Mil	lls OH (la	t 39 54 3	ln long 080 5	5 27w)
100	20	270	240	30	20	10	10	
3113550	D-2	McMahon (	C at Boll:	ire OH (	lat 40 00	39n long	080 45 45w)	
70	140	490	470	20	90	0	90	
03114250	D-3	Sunfish (	C at Came	on OH (1a	at 39 46 0	On long O	80 56 09w)	
40	30	150	120	30	20	0	20	
4000230805320	00 D-4		C at Glend	oe OH (1	at 40 00 2	3n long 0	80 53 20w)	
40	50	220	140	80	20	0	20	
0001308053300	0 D-5	Williams	C at Glend	CO OH (la	40 00 13	n long 08	0 53 30w)	
40	10	150	110	40	<10		10	
4002250805041	00 D-6	L McMaho	n C nr Ne	ffs Oh (1	at 40 02 2	5n long 0	80 50 41w)	
3,370	130	3,100	3,100	20	120	10	110	
3954190810448	00 D-7				OH (lat 39		ong 081 04 48	w)
140	30	380	350	30	30	20	10	
3954440810250	00 D-8					44n long	081 02 50w)	
80	30	240	160	80	50	20	30	
3955020805757	00 D-9	Bend Fk	nr Armstr	ongs Mill:	s OH (lat	39 55 02n	long 080 57	57w)
0	30	120	100	20	<10		10	
3953330805413	000 D-10	Pea Vine	C nr Arms	trongs Mi	lls OH (la	t 39 53 3	3n long 080 5	4 13w)
30	10	120	100	20	20	10	10	
3947120810701	.00 D-11	Sunfish	C nr Wood	dsfield O	H (1at 39		ng 081 07 01w	
10	30	150	100	50	40	0	40	
3948270810653	00 D-12	Baker F	k nr Wood	sfield OH	(lat 39 4	8 27n lon	g 081 06 53w)	
20	20	150	110	40	20	0	20	
30								
	00 D-13	Piney F	k nr Wood	sfield OH	(lat 39 4	6 45n lon	g 081 00 41w)	

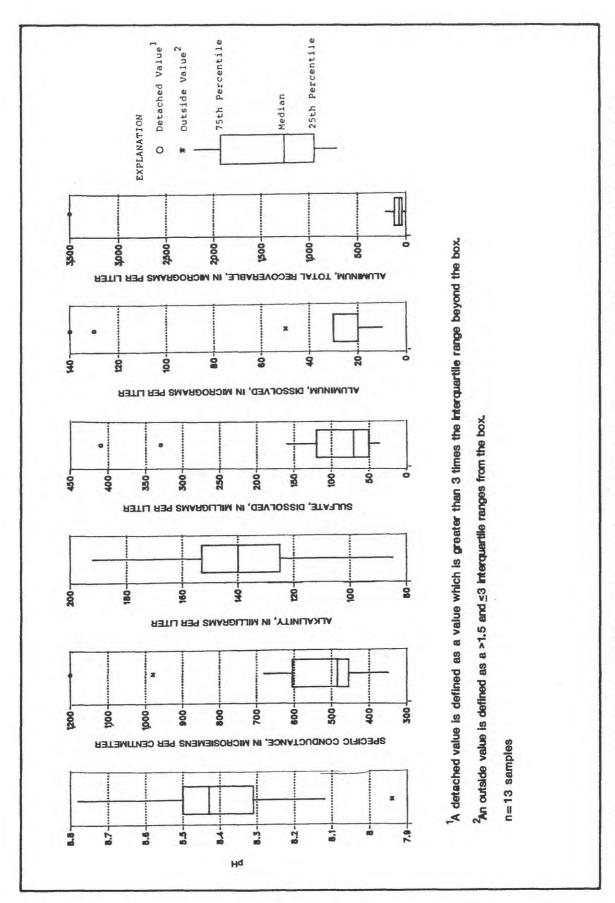


Figure 21.—Box plots showing the range, percentiles, and median values of constituents at surface-water sites in the McMahon, Captina, and Sunfish Creeks basin.

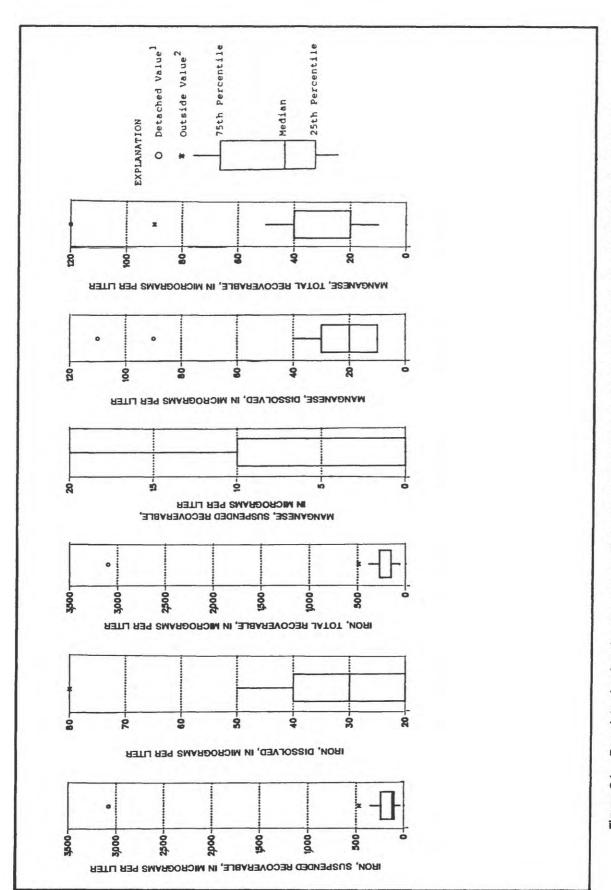


Figure 21.—Box plots showing the range, percentiles, and median values of constituents at surface-water sites in the McMahon, Captina, and Sunfish Creeks basin--Continued.

Specific conductance ranged from 350  $\mu$ S/cm (at sites D-12 and D-13) to 1,200  $\mu$ S/cm (at site D-6), with a median of 485  $\mu$ S/cm.

pH ranged from 7.94 (at site D-6) to 8.78 (at site D-11), with a median of 8.43. Alkalinity ranged from 85 mg/L as CaCO<sub>3</sub> (at site D-12) to 192 mg/L (at site D-6), with a median of 140 mg/L.

Dissolved-sulfate concentrations ranged from 37 mg/L (at site D-7) to 410 mg/L (at site D-6), with a median of 71 mg/L.

Total-iron concentrations ranged from 60  $\mu$ g/L (at site D-13) to 3,100  $\mu$ g/L (at site D-6), with a median of 150  $\mu$ g/L. Dissolved-iron concentrations ranged from 20  $\mu$ g/L (at sites D-2, D-6, D-9, D-10, and D-13) to 80  $\mu$ g/L (at sites D-4 and D-8), with a median of 30  $\mu$ g/L.

Concentrations of total aluminum ranged from 10  $\mu g/L$  (at site D-13) to 3,500  $\mu g/L$  (at site D-6), with a median of 70  $\mu g/L$ . Dissolved-aluminum concentrations ranged from 10  $\mu g/L$  (at sites D-5 and D-10) to 140  $\mu g/L$  (at site D-2), with a median of 30  $\mu g/L$ .

Total-manganese concentrations ranged from less than 10  $\mu g/L$  (at sites D-5, D-9, and D-13) to 120  $\mu g/L$  (at site D-6), with a median of 20  $\mu g/L$ . Dissolved-manganese concentrations ranged from 10  $\mu g/L$  (at sites D-1, D-5, D-7, D-9, D-10, and D-13) to 110  $\mu g/L$  (at site D-6), with a median of 20  $\mu g/L$ .

## SUMMARY

Twenty basins were selected in the coal-bearing region of Ohio for the study of surface- and ground-water quality. The study area is mostly within the unglaciated eastern part of Ohio with local relief as much as 500 feet. The area is underlain by limestones, sandstones, shale, and coal of Mississippian, Pennsylvanian, and Permian age that dip to the southeast at approximately 30 feet per mile. Unconsolidated deposits are found chiefly in old valleys that now have no major drainage. These deposits are chiefly glacial outwash, and some have a yield greater than the 5- to 10-gal/min yield of some in bedrock aquifers.

One to three long-term surface-water sites were selected in each basin for sampling three times during low flow. All samples were analyzed for specific conductance, pH, alkalinity, acidity, dissolved sulfate, total and dissolved iron, total and dissolved manganese, and total and dissolved aluminum. Specific conductance in these 100 samples ranged from 270 µS/cm to 2,800 µS/cm.

Dissolved-sulfate concentrations ranged from 17 mg/L to 1,400 mg/L. Dissolved-iron concentrations ranged from less than 10  $\mu$ g/L to 54,000  $\mu$ g/L, dissolved-manganese concentrations ranged from 10  $\mu$ g/L to 16,000  $\mu$ g/L, and dissolved-aluminum concentrations ranged from less than 10  $\mu$ g/L to 18,000  $\mu$ g/L. Elevated concentrations of these constituents commonly are used as indicators of acid mine drainage.

Five individual basins selected for intensive investigation were Stillwater Creek basin; Symmes, Ice, and Indian Guyan Creeks basin; Moxahala Creek basin; Little Beaver Creek basin; and McMahon, Captina, and Sunfish Creeks basin. Intensive study included one-time sampling of 12 to 13 surface-water sites within the basin, study of a productive aquifer in parts of the three basins, and a description of the geologic setting. If a productive aquifer were present, three to five ground-water samples were collected, and water levels associated with that aquifer were measured. In addition to all of the physical properties and chemical constituents analyzed for in the stream samples, the ground-water samples were analyzed for silica, calcium, magnesium, sodium, chloride, potassium, total organic carbon, and total dissolved solids.

The Stillwater Creek basin is underlain by shale, sandstone, coal, and limestone of the Conemaugh Formation. Eleven surface-water samples had a range in specific conductance of 390 to 3,300  $\mu\text{S/cm}$ . Dissolved-sulfate concentrations ranged from 15 to 1,900 mg/L. Concentrations of dissolved iron ranged from less than 10 to 130  $\mu\text{g/L}$ , dissolved-manganese concentrations ranged from 90 to 1,700  $\mu\text{g/L}$ , and dissolved-aluminum concentrations ranged from less than 100 to 300  $\mu\text{g/L}$ .

Ground-water quality was investigated at two locations within the Stillwater Creek basin. At the Village of Flushing, as much as 200 gal/min have been pumped from wells in the glacial outwash deposits along Jockey Hollow Run. The water is a very hard calcium sulfate type that is extremely high (883 to 1,200 mg/L) in dissolved solids, which may indicate acid mine drainage. Dissolved-sulfate concentrations ranged from 390 to 590 mg/L, and exceeded the concentrations found in other typical Ohio outwash aquifers and the OEPA standard for a public supply. Elevated concentrations of dissolved iron (17 to 3,300 µg/L) and manganese (780 to 3,100 µg/L) also were found. Samples also were collected from Tappan Lake Park near Deersville. The wells in the park are drilled into the Conemaugh Formation, which yields 5 to 10 gal/min. Analyses show that Tappan Park has a moderately hard to hard calcium bicarbonate type ground water. Dissolved-solids concentrations ranged from 194 to 311 mg/L. Concentrations of dissolved iron (710 to 1,200 µg/L) and total manganese (80 to 370 µg/L) exceeded OEPA standards in all samples.

Strata beneath the Symmes, Ice, and Indian Guyan Creeks basin range from sandstone and shale of the Allegheny Formation in the western part of the basin to the predominantly shaly rock of the Conemaugh and Monongahela Formations to the east. Alluvial deposits of gravel, sand, silt, and clay overlie the bedrock throughout the uplands and generally are less than 50 feet thick. Most bedrock wells yield less than 5 gal/min. Specific conductance of 12 surface-water samples ranged from 250 to 465 µS/cm. Dissolved-sulfate concentrations ranged from 53 to 170 mg/L. Concentrations of dissolved iron ranged from 20 to 150 µg/L, and dissolved-manganese concentrations ranged from 120 to 1,400 µg/L.

Bedrock in the Moxahala Creek basin consists of sandstone, shale, and limestone of Mississippian and Pennsylvanian age. Glacial deposits of clay, silt, sand, and gravel overlie the bedrock in the northwestern part of the basin. Yields of 5 to 10 gal/min are typical of bedrock wells. Wells in glacial deposits have yielded as much as 200 gal/min, but 20 gal/min is typical. Specific conductance of 13 surface-water samples ranged from 520 to 4,600 µS/cm. Dissolved-sulfate concentrations ranged from 42 to 2,800 mg/L. Dissolved-iron concentrations ranged from 30 to 120,000 µg/L, dissolved-manganese concentrations ranged from 80 to 23,000 µg/L, and dissolved-aluminum concentrations ranged from 10 to 24,000 µg/L.

Ground-water quality was investigated in the glacial outwash deposits of the Jonathan Creek Valley near Glenford. This aquifer has a hard to very hard calcium bicarbonate type water that is moderately high in dissolved solids. Most coal mining is located to the east and southeast of Glenford, and probably has little effect on the water quality in this area. Dissolved-solids concentrations ranged from 268 to 384 mg/L. Hardness ranged from 220 to 330 mg/L as CaCO3. Three of the five samples exceeded the OEPA limits for these constituents in a public water supply. Concentrations of dissolved iron and dissolved manganese ranged from less than 3 to 4,400  $\mu \text{g/L}$  and less than 1 to 290  $\mu \text{g/L}$ , respectively.

Rock formations underlying the Little Beaver Creek basin consist of sandstone, limestone, coal, and shale of the Allegheny and Conemaugh Formations. The southern part of the basin is unglaciated and has the potential to provide yields of less than 5 gal/min to wells. Glacial till and outwash deposits are located along Little Beaver Creek and its tributaries in the northern part of the basin and may yield 5 to 25 gal/min to wells. Specific conductance of 12 surface-water samples ranged from 330 to 1,210  $\mu\text{S/cm}$ . Dissolved-sulfate concentrations ranged from 41 to 250 mg/L. Dissolved-iron concentrations ranged from 20 to 80  $\mu\text{g/L}$ , and dissolved-manganese concentrations ranged from 20 to 760  $\mu\text{g/L}$ .

Ground-water quality near Lisbon, mainly in the shallow sandstone of the Conemaugh Formation, is a very hard calcium bicarbonate type water. Mining is confined to small operations scattered throughout the basin, and one sample (CO-1) downgradient from such a mine showed evidence of possible acid mine drainage. Specific conductance of five samples ranged from 511 to 1,300  $\mu$ S/cm. Hardness ranged from 260 to 790 mg/L. Total organic carbon ranged from 0.4 to 0.8 mg/L in four wells, but was elevated to 6.0 mg/L in a fifth well. Concentrations of dissolved iron ranged from 13 to 2,500  $\mu$ g/L, and three samples exceeded the OEPA standard for a public supply. Total-manganese concentrations (60 to 3,200  $\mu$ g/L) exceeded the standard in all samples.

The McMahon, Captina, and Sunfish Creeks basin is underlain by rocks of the Conemaugh and Monongahela Formations of Pennsylvanian age and the Dunkard Group of Pennsylvanian and Permian age. Yields of 2 to 5 gal/min are developed in the bedrock. Alluvial deposits on the creek flood plains may yield up to 25 gal/min to wells. Specific conductance of 13 surface-water samples ranged from 350 to 1,200  $\mu\text{S/cm}$ . Dissolved-sulfate concentrations ranged from 37 to 410 mg/L. Dissolved-iron concentrations ranged from 20 to 80  $\mu\text{g/L}$ , and dissolved-manganese concentrations ranged from 10 to 110  $\mu\text{g/L}$ .

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# Table 2 .-- Water-quality data for long-ters surface-water sites

The following table contains water-quality data collected at long-term sites. The sites are listed alphanumerically by basin-identification number. For each site, the station number, local identifier, station name, latitude, and longitude are given. Data for selected properties and constituents are listed for each observation. Missing data are represented with a dash. Site information is presented in the following format:

	Basin iden-			
Station number	tifi- cation	Station name	Latitude	Longitude
385612082095900	J-1 L K	YGER C NR KYGER OH	(LAT 385612	LONG 0820959)

#### STANDARD ABBREVIATIONS USED IN STATION NAMES

AB	Above	F	Fork	NR	Near
В	Branch	G	Great	R	River
BK	Brook	L	Little	RN	Run
BL	Below	LK	Lake	S	South
C	Creek	M	Middle	TR	Tributary
E	East	N	North	W	West

## OTHER SYMBOLS AND ABBREVIATIONS USED IN TABLE 2

UG/L	micrograms per	liter	CL	chloride (C1)
UG/G	micrograms per	gram	NI	nickel (Ni)
US/CM	microsiemens p	er centimeter	504	sulfate (SOA)
FE	iron (Fe	)	CACO3	calcium carbonate (CaCO3)
AS	arsenic (As	)	HCO3	bicarbonate (HCO3)
HG	mercury (Hg	)	CFS	cubic feet per second
CU	copper (Cu	)	FM	from
CR	chromium (Cr	)	RECOV.	recoverable
AL	aluminum (Al	.)	FET-FLD	fixed endpoint titration -
MN	manganese (Mn	)		field determination
ZN	zinc (Zn	)		
MG/L	milligrams per	liter		

DATE	STREAM- FLOW, INSTAN- TANEOUS (CFS)	CON-	PH (STAND- ARD UNITS)	TEMPER- ATURE (DEG. C)	ACIDITY (MG/L AS CACO3)	FIELD MG/L AS	DIS- SOLVED (MG/L	RECOV- ERABLE (UG/L
03109500	A-1	L BEAVER	C NR EAS	ST LIVERPO	OL OH (LA	T 40 40	33N LONG	080 32 27W)
OCT 1985								
02 JUN 1986	61	950	8.20	12.0		193	230	200
25	70	750	8.40	19.5		105	170	200
22	93	700	8.83	12.5		130	160	90
		A-2 M F I				T 40 43	22N LONG	080 38 03w)
OCT 1986 21		765	8.81	12.0		147	150	130
03110000		B-1 YELLOW	C NR H	MMONDSVIL	LE OH (LA	T 40 32	16N LONG	080 43 31W)
OCT 1985	6.2	770		12.5		155		200
JUN 1986 25	24	497	8.40	21.5		93	120	100
OCT 22	14	570	8.42	13.5		101	150	80
401857080	391700	B-2 CROSS	C NR MIN	GO JUNCTIO	N OH (LA	r 40 18 5	7N LONG	080 39 17W)
OCT 1985								
	13	1,750	7.90	12.5		173	780	500
25	16	1,420	8.70	22.5		175	610	510
20	20	1,300	8.78	10.5		135	600	840
401716080	451300	B-3 MCINT	YRE C NE	SMITHFIEL	D OH (LA	r 40 17 1	6N LONG	080 45 13W)
OCT 1986 20		2,000				203		60
03111548		C-1 WHEELI	NG C BL	BLAINE OR	(LAT 40 (	04 Oln LO	NG 080 4	8 31w)
OCT 1985		2,400					1,000	
JUN 1986						107	0.00	450
OCT		2,400					960	
20	22	2,250	8.24	7.5		220	1,100	420
03111500		C-2 SHORT	C NR DI	LLONVALE (	H (LAT 4)	11 36N	LONG 080	44 04W)
OCT 1985	24	2,600	9.00	12.5		390	1,200	400
02 JUN 1986								
25 OCT	31	2,360	8.40	22.5		198	1,100	310
20				8.5				320
03114000	D-1	CAPTINA C	AT ARMST	RONGS MILI	S OH (LA	r 39 54 3	IN LONG	080 55 27W)
OCT 1985							450	100
02 JUN 1986	1.7							
24 OCT	5.7			24.5 8.5		137		160
03113550		D-2 MCMAI	ON C AT	BELLAIRE (	OH (LAT 4)	U 00 39N	LONG U80	45 45W)
OCT 1985 03	3.2	1,700	7.50	12.5		151	730	100
JUN 1986 24							360	290
OCT	15		8.31			146	330	210

LONG-TERM SURFACE-WATER SITES

ALUM- INUM, SUS-	ALUM-	IRON,	IRON, SUS-			MANGA- NESE,	MANGA- NESE, SUS-	MANGA-	
PENDED	INUM,	TOTAL	PENDED	IRON,		TOTAL	PENDED	NESE,	
RECOV- ERABLE	DIS- SOLVED	RECOV-	RECOV- ERABLE	DIS- SOLVED		RECOV- ERABLE			
(µg/L	(PG/L	(PG/L	(UG/L	(JG/L		(PG/L	(µG/L	(DG/L	
AS AL)	AS AL)	AS FE)	AS FE)	AS FE)		AS MN)	AS MN)	AS MN)	
03109500	A-1 L B	EAVER C N	R EAST LI	VERPOOL	OH	(LAT 40	40 33N	LONG 080	32 27W)
	<100	340	330	10		50	30	20	
170	30	290	240	50		130	60	70	
80	10		-	50		50		30	
03109100			EAVER C NR						38 03W)
110	20	560	530	30		80	40	40	
03110000	B-1	YELLOW C	NR HAMMON	DSVILLE	ОН	(LAT 40	32 16N	LONG 080	43 31W)
	<100	370		10			10	20	
40	60	380	350	30		40	20	20	
50	30			20			10	10	
401857080391	700 B-2	CROSS C N	R MINGO J	UNCTION	OH	(LAT 40	18 57N	LONG 080	39 17W)
300	200	310		<10		220	50	170	
280	230	520	470	50		240	30	210	
770	70	560	530	30		170	40	130	
401716080451	300 B-3	MCINTYRI	C NR SMI	THFIELD	ОН	(LAT 40	17 16N	LONG 080	45 13W)
7.7	40		170				0	100	
03111548	C-1	WHEELING	C BL BLA	INE OH	LAT	40 04 0	IN LONG	080 48	31W)
	<100	490	460	30		650	140	510	
300	150	1,600	1,600	20		140	10	130	
320	100	920	900	20		190	0	190	
03111500			NR DILLON	VALE OH	(LA	T 40 11	36N LON	G 080 44	04W)
200	200	800	760	40		90	50	40	
70	240	590	570	20		60	20	40	
210	110	700	680	20		90	10	80	
03114000	D-1 CAPT	PINA C AT	ARMSTRONG	S MILLS	он	(LAT 39	54 31N	LONG 080	55 27W)
0	100	240	200	40		40	10	30	
130	30	40	0			70	50		
100		270				20			
03113550	p-2	MCMAHON	C AT BELL	AIRE OH	(L	T 40 00	39N LON	G 080 45	45W)
0	100	550	530	20		70	10	60	
210	80	950	910	40		120	20	100	
70	140	490	470	20		90	0	90	

DATE	TANEOUS	CON- DUC-	ARD	ATURE		LINITY, FIELD (MG/L AS		RECOV- ERABLE (UG/L
03114250		D-3 SUNF	ISH C AT	CAMERON (	OH (LAT 39	46 00N	LONG 080	56 09W)
OCT 1986 22			8.31	10.5		125	52	70
03117500		E-1 SAND	Y C AT W	AYNESBURG	OH (LAT 4	0 40 21N	LONG 081	15 36W)
OCT 1985 02 JUN 1986	35	1,660						200
24	70	600	8.70	21.5		117	110	100
OCT 22	68	340	8.03	13.5		116	120	70
403823081	1213700 E	E-2 NIMISH	ILLEN C A	T SANDYVI	LLE OH (L	AT 40 38	23 LONG	081 21 37)
OCT 1985								
03	70	1,440	7.60	12.0		207	210	<200
JUN 1986 24		1,410	8.50	22.0		240	180	140
OCT 22	81	1,120	8.12	15.0		216	180	100
403426081								81 21 19W)
OCT 1985 03 JUL 1986	12	665		11.5		88	180	200
07 OCT	202	340	7.70	24.0		123	65	2,700
22	46	310	7.66	14.0		79	120	250
		R C AB BEA	CH CITY	DAM AT BE	ACH CITY	H (LAT	03924 LO	NG 0813437
OCT 1985 02		755						800
JUN 1986			7.00	23.5		222	66	1,000
24 OCT	19	650						
22		720		18.5			74	
		G-2 NEWM	N C NR M	ASSILLON	OH (LAT 4	49 22N	LONG 081	
OCT 1986 22	1.8	460	8.24	13.0		209	110	200
03127500	н-1	STILLWATE	ER C AT U	HRICHSVIL	LE OH (LAT	r 40 23 1	ON LONG	081 20 50W)
SEP 1985	12	1,200	7.60	19.0		206	390	700
JUN 1986 24	65	820	7.60	22.5		137	270	480
22	88	890	7.69	12.5				300
03129100		I-1 WHITE	E EYES C	NR FRESNO	OH (LAT	40 18 171	LONG DE	1 45 01W)
OCT 1985								
	. 0.79	445	6.60	14.0	:	95	63	100
	14	380	7.80	24.5		96	58	330
23	14	420	7.67	12.5		89	66	240

ALUM- INUM, SUS- PENDED RECOV- ERABLE (µG/L AS AL)	ALUM- INUM, DIS- SOLVED (PG/L AS AL)	RECOV- ERABLE		IRON, DIS- SOLVED (µG/L AS FE)	(µg/L	MANGA- NESE, SUS- PENDED RECOV- ERABLE (PG/L AS MN)	MANGA- NESE, DIS- SOLVED (PG/L AS MN)	
03114250	D-3	SUNFISH	C AT CAME	RON OH (	LAT 39 46	00N LONG	080 56 09	()
40	30	150	120	30	20	0	20	
03117500			AT WAYNES		(LAT 40 40	21N LONG	081 15 36	5W)
		270		30	80	0	90	
60	40	750	730	20	400	60	340	
50	20	520	500	20	410	10	400	
403823081213	700 E-2 N	IMISHILLI	EN C AT SA	NDYVILLE	OH (LAT 4	0 38 23 L	ONG 081 21	37)
	200	520	500	20	110	20	90	
100	40	790	710	80	100	60	40	
90	10	840	810	30	170	0	170	
403426081211	900 F-1	CONOTTO	N C NR SOM	ERDALE O	H (LAT 40	34 26N LO	NG 081 21	19W)
	<100	1,600	1,500	90	1,900	0	1,900	
2,670	30	8,900	8,800	50	1,400	530	870	
200	50	2,000	1,700	330	1,000	0	1,100	
03123000 G-1	SUGAR C	AB BEACH	CITY DAM	AT BEACH	CITY OH	LAT 40392	4 LONG 081	3437)
700			1,600				170	
980	20	2,300	2,300	20	370	50	320	
690	180	1,700	1,320	380	220	60	160	
03116950	G-2	NEWMAN (	C NR MASSI	LLON OH	(LAT 40 49	22N LONG	081 33 06	5W)
180	20	940	910	30	180	30	150	
03127500	H-1 STI	LLWATER	C AT UHRIC	HSVILLE	OH (LAT 4	23 10N I	ONG 081 20	5 OW
400	300	1,700		<10	1,300	900	400	
450	30	1,100		<10		60	540	
	40	840	820	20	450	30	420	
03129100	I-1	WHITE E	YES C NR F	RESNO OH	(LAT 40		G 081 45	01W)
	<100		2,200			30	460	
	17.00							
250	80	1,500	1,400	100	310	30	280	

DATE	STREAM- FLOW, INSTAN- TANEOUS (CFS)	CON- DUC- TANCE	(STAND-	TEMPER- ATURE	(MG/L AS	LINITY, FIELD (MG/L AS	SOLVED	RECOV- ERABLE (UG/L
40162408	1363400							081 36 34W)
OCT 1985								
02 JUN 1986		710		10.0		148	120	100
OCT				20.5		67	80	110
				12.0			76	160
				ностои он				
OCT 1985								
01 JUN 1986	0.43	470				125	30	277
23 OCT	6.1			24.0		128	42	250
				10.0			54	80
				WARSAW OH				
SEP 1985		1 070	6 00	16.0		124	430	100
30 JUN 1986	0.19	1,070	0.80	23.0				410
							430	
21				8.0			250	
	1323000	K-1 WILLS	C AT PL	EASANT CITS	OH (LA	39 54 1	7N LONG C	81 32 30W)
OCT 1985								
01 JUN 1986				15.0				
25 OCT	4.7	2,800	7.90	21.5		200	1,200	680
22	11	2,050	8.21	11.5		201	930	520
40011708	362600	L-1 CROOK	ED C NR	CAMBRIDGE (	H (LAT	0 01 17N	LONG 081	36 26W)
OCT 1985								
02 JUN 1986				13.5		169	140	200
25 OCT	1.0	640	7.70	21.5		145	110	620
	4.6	630	7.97	12.5		149	110	240
					OH (LAT	40 09 2	ON LONG O	81 43 29w)
ост 1985								
01 JUN 1986	1.3	960	6.70	14.0		52	410	300
26 OCT	4.5	660	8.30	20.5		79	200	180
23	17	700	7.44	13.0		85	240	110
41061608	2075500	4-1 WAKATOM	IKA C NR	FRAZEYSBU	RG OH (L	AT 41 06	16N LONG	082 07 55W)
SEP 1985								
30 JUN 1986	5.5	445	6.80	15.0		131	17	200
26 OCT	25	410	7.70	18.0		81	21	540
21	41	410	7.50	8.5		75	25	60

ALUM- INUM, SUS- PENDED RECOV- ERABLE (µG/L AS AL)	ALUM- INUM, DIS- SOLVED (µG/L AS AL)		IRON, SUS- PENDED RECOV- ERABLE (µG/L AS FE)	(µG/L	ERABLE	RECOV-	NESE, DIS- SOLVED (UG/L	
401624081363	1400 I-2 B	UCKHORN C	AT NEWCO	MERSTOWN	OH (LAT 40	16 24N	LONG 081	36 34W)
	<100	230	170	60	60	20	40	
70	40	570	500	70	330	30	300	
140	20	770	750	20	550	0	560	
03140000	J-1	MILL C N	R COSHOCT	ON OH (LA	T 40 21 461	LONG (	81 51 45	v)
	<100	2,700	2,500	190	250	140	110	
220	30	1,200	1,100	70	240	40	200	
	<10	1,600	1,500	60	270	0	280	
401936082001					AT 40 19 36			w)
	<100	220	210	10	130	50	80	
390	20	1,200	1,200	20	600	140	460	
30	10	420	380	40	280	0	280	
395417081323		WILLS C	AT PLEASAN	T CITY OF	H (LAT 39 5	4 17N L	ONG 081 3	2 30W)
300					240			
640	40	950	940	10	350	180	170	
490	30.	610	570	40	170	90	80	
400117081362					(LAT 40 01			26W)
	<100	380	350	30	450	60	390	
600	20	1,200	1,200	30	860	70	790	
220	20	630		<10	290	200	90	
400920081432					H (LAT 40 0			3 29W)
100	200	1,600	1,400	160	1,700	0	1,700	
150	30	1,100	1,000	70	800	0	800	
100	10	890	790	100	940	0	970	
410616082075	500 M-1 W	AKA TOMIKA	C NR FRAZ	EYSBURG (	OH (LAT 41	06 16N	LONG 082	07 55W)
	<100	540	460	80	150	30	120	
520	20	1,500	1,400	110	200	70	130	
50	10	840	750	90	90	0	90	

DATE	FLOW, INSTAN- TANEOUS	SPE- CIFIC CON- DUC- TANCE (US/CM)	46444	********		FIELD (MG/L AS	10001 0	RECOV- ERABLE (UG/L
		M-2 L WAKA						
SEP 1985								
30 JUN 1986		1,690						200
OCT		1,400					590	
		970				104	390	80
03149500		N-1 SALT	C NR CHAI	NDLERSVIL	LE OH (LA	т 39 54 3	IN LONG	081 51 38W)
OCT 1985							52	200
01 JUN 1986		580		15.0				
OCT		510						140
		480		7.5		110		140
		N-2 MEIGS	C NR BEY		(LAT 39 3	OON LON	G 081 42	
SEP 1985	2.6	650	8.00	17.5				
JUN 1986 23	19	1,200	7.70	26.5		145	490	380
OCT 21	7.4	1,400	7.70	9.0		163	630	270
		O-1 MOXAH						182 01 11W)
		U-I HOXAN						
OCT 1985 01 JUN 1986	19	1,800						
23 OCT	41	1,300	6.50	27.5		16	530	320
22	25	1,600	6.30	12.0	9.0	25	800	360
		O-2 MOXAH			OH (LAT 39	51 17N	LONG 082	03 23W)
OCT 1986 22		2,600	3.44	11.0	148		1,400	8,600
39521408	2054700	-3 JONATHA	N C AT WE					082 05 47W
OCT 1986 22		1,050	7.46				310	280
03156700		P-1 RUSH	C NR SUG	R GROVE	OH (LAT 3	9 38 18N	LONG 082	30 42W)
SEP 1985	18	890	7.60	14.0		118	280	100
JUN 1986 23				26.0		85	230	340
OCT 21	24	515	8.20	7.5		101	120	340
03157000		P-2 CLEAR	C NR RO	CKBRIDGE	OH (LAT 3	9 35 18N	LONG 082	34 43W)
	27	430						40
03158200		Q-1 MONDA	Y C AT DO	DANVILLE	OH (LAT 3	9 26 07N	LONG 082	11 30W)
SEP 1985		1,200						18,000
JUN 1986		1,200				0		8,400
OCT 21		1,150				0	480	7,800

SUS-		4000	IRON,		MANGA-	MANGA- NESE,		
	ALUM-	IRON,	SUS-	****	NESE,	SUS-	MANGA-	
	INUM,	TOTAL	PENDED		TOTAL RECOV-		NESE,	
RECOV-	DIS-		RECOV-		5177777			
	SOLVED		ERABLE	SOLVED		ERABLE		)
(µG/L AS AL)	(JG/L AS AL)	(µG/L AS FE)	AS FE)	(µG/L AS FE)		(µG/L AS MN)	AS MN)	
400912082014			MIKA C NR	TRINWAY	OH (LAT 40	09 12N I	ONG 082	01 47W
100	100	570	540	30	940	80	860	
130	30	610	570	40	830	0	830	
70	10	710	650	60	1,100	0	1,100	
03149500		SALT C	NR CHANDL	ERSVILLE	OH (LAT 39	54 31N I	ONG 081	51 38W
	<100	740	720	20	440	0	510	
120	20	790	680	110	370	50	320	
120	20	980	850	130	260	0	260	
03150250					т 39 36 00			)
	<100	540	460	80	400	20	380	
350	30	650	590	60	120	90	30	
	<10	520	500	20	140	60	80	
395337082011					OH (LAT 39			01 11W
0	8,100	1,200	200	1,000	15,000	2,000	13,000	
280	40		180	30	6,600	0	6,800	
			250		6,400		6,600	
								-
03148400	0-2	MOXAHAL	A C AT RO	BERTS OH	(LAT 39 51	17N LONG	082 03	23W)
03148400	0-2	MOXAHAL			14,000			23W)
400	0-2 8,200 700 0-3 J	MOXAHAL 11,000 ONATHAN	0 C AT WHITI	11,000	14,000	0 9 52 14N	16,000	
400	0-2 8,200 700 0-3 J	MOXAHAL 11,000 ONATHAN	0 C AT WHIT	11,000 E COTTAGE	14,000 OH (LAT 3	0 9 52 14N	16,000 LONG 082	
400 400 3952140820547 200	0-2 8,200 700 0-3 J	MOXAHAL 11,000 ONATHAN 420	O C AT WHITE	11,000 E COTTAGE	14,000 OH (LAT 3	0 9 <b>52 14N</b> 0	16,000 LONG 082 2,600	05 47
400 400 3952140820547 200 03156700	0-2 8,200 700 0-3 J 80	MOXAHAL 11,000 ONATHAN 420 RUSH C	0 C AT WHITH 390 NR SUGAR (	11,000 E COTTAGE 30 GROVE OH	14,000 OH (LAT 3	0 9 52 14N 0 18N LONG	16,000 LONG 082 2,600	05 47
400 400 3952140820547 200 03156700	0-2 8,200 700 0-3 J 80	MOXAHAL 11,000 ONATHAN 420 RUSH C	O C AT WHITE 390 NR SUGAR (	11,000 E COTTAGE 30 GROVE OH	14,000 OH (LAT 3 2,500 (LAT 39 38	0 9 52 14N 0 18N LONG	16,000 LONG 082 2,600 3 082 30 2,100	05 47
03148400 400 3952140820547 200 03156700	0-2 8,200 700 0-3 J 80 P-1 100 30	11,000 ONATHAN 420 RUSH C 350 870	O C AT WHITE 390 NR SUGAR (	11,000 E COTTAGE 30 GROVE OH	14,000 OH (LAT 3 2,500 (LAT 39 38 2,100 2,000	0 9 52 14N 0 18N LONG	16,000 LONG 082 2,600 3 082 30 2,100	05 47
03148400 400 3952140820547 200 03156700 0 310 320	0-2 8,200 700 0-3 J 80 P-1 100 30 20	11,000 ONATHAN 420 RUSH C 350 870 1,100	390 NR SUGAR ( 330 790 1,100	11,000 E COTTAGE 30 GROVE OH 20 80 30	14,000 OH (LAT 3 2,500 (LAT 39 38 2,100 2,000	0 9 52 14N 0 18N LONG 0 100	16,000 LONG 082 2,600 082 30 2,100 1,900 1,700	05 471 42w)
03148400 400 3952140820547 200 03156700 0 310 320	0-2 8,200 700 0-3 J 80 P-1 100 30 20	MOXAHAL  11,000  ONATHAN  420  RUSH C  350  870  1,100  CLEAR C	390 NR SUGAR ( 330 790 1,100	11,000 E COTTAGE 30 GROVE OH 20 80 30	14,000 OH (LAT 3 2,500 (LAT 39 38 2,100 2,000 1,700 (LAT 39 35	0 9 52 14N 0 18N LONG 0 100	16,000 LONG 082 2,600 082 30 2,100 1,900 1,700	05 47 42w)
03148400 400 3952140820547 200 03156700 0 310 320	0-2 8,200 700 0-3 J 80 P-1 100 30 20 P-2	MOXAHAL  11,000  ONATHAN  420  RUSH C  350  870  1,100  CLEAR C	390  NR SUGAR (  330  790  1,100  NR ROCKE	11,000 E COTTAGE  30 GROVE OH  20 80 30 RIDGE OH	14,000 OH (LAT 3 2,500 (LAT 39 38 2,100 2,000 1,700 (LAT 39 35	0 9 52 14N 0 18N LONG 0 100 0	16,000 LONG 082 2,600 082 30 2,100 1,900 1,700 6 082 34 40	05 471 42W)
03148400 400 3952140820547 200 03156700 0 310 320 03157000 20	0-2 8,200 700 0-3 J 80 P-1 100 30 20 P-2 20	11,000 ONATHAN 420 RUSH C 350 870 1,100 CLEAR C 280	390  NR SUGAR (  330  790  1,100  NR ROCKB!	11,000 E COTTAGE  30 GROVE OH  20 80 30 RIDGE OH  50	14,000 OH (LAT 3 2,500 (LAT 39 38 2,100 2,000 1,700 (LAT 39 35	0 9 52 14N 0 18N LONG 0 100 0 18N LONG	16,000 LONG 082 2,600 3 082 30 2,100 1,900 1,700 6 082 34 40 3 082 11	05 47 42w)
03148400 400 3952140820547 200 03156700 0 310 320 03157000 20 03158200	0-2 8,200 700 0-3 J 80 P-1 100 30 20 P-2 20	MOXAHAL  11,000  ONATHAN  420  RUSH C  350  870  1,100  CLEAR C  280  MONDAY  4,100	0 C AT WHITE 390 NR SUGAR ( 330 790 1,100 NR ROCKHI 230 C AT DOAN	11,000 E COTTAGE  30 GROVE OH  20 80 30 RIDGE OH  50 VILLE OH	14,000 OH (LAT 3 2,500 (LAT 39 38 2,100 2,000 1,700 (LAT 39 35 40 (LAT 39 26	0 9 52 14N 0 18N LONG 0 100 0 18N LONG 0 07N LONG	16,000 LONG 082 2,600 3 082 30 2,100 1,900 1,700 6 082 34 40 3 082 11	05 47 42w)

	STREAM- FLOW, INSTAN- TANEOUS (CFS)	CI CO DU TA	N- C-	PH (STAND- ARD UNITS)	TEMPER- ATURE (DEG. C)	(MG/L AS	ALKA- LINITY, FIELD (MG/L AS CACO3)	DIS- SOLVED (MG/L	RECOV- ERABLE
							23 42N L		
OCT 1986 22					8.0			1,000	3,500
			RACCO	ON C AT	VINTON OH	(LAT 38	58 26N LOI	NG 082 20	18W)
SEP 1985	3.9				15.0		20	150	200
JUN 1986 25	28		390	6.70	23.0		19	110	110
22	11		450	7.13	9.0		33	140	70
39094108	2212200	R-2	ELK F	K NR RAI	CLIFF OH	(LAT 39 C	9 41N LONG	082 21	22W)
OCT 1986 22				7.35			48	130	110
03201988		S-1	L RAC	COON C N	R VINTON	OH (LAT 3	8 57 11N I	ONG 082	21 56W)
SEP 1985	8.7				14.0	1979797	8		1,000
JUN 1986 25	4.9		455	6.80	20.0		6	160	290
	6.6		790	4.75	8.5	31	0	340	4,100
03160105		S-2	CAMPA	IGN C NE	GALLIPOL	IS OH (LA	T 38 53 51	N LONG O	82 11 31W
OCT 1986 23					12.5		23		100
383005082	2280600	T-1	SYMME	S C NR G	ETAWAY OH	(LAT 38	30 05N LON	G 082 28	06W)
OCT 1985 01	1.3		410		15.0		105		200
24			270	7.00	25.0		73	42	1,400
	13		405	7.68	10.0		75	100	140
							T 38 27 15		
OT 1985							182		100
24	2.2		565	7.60	23.0		86	150	310
23	2.6		615	7.76	10.5		85	200	160

LONG-TERM SURFACE-WATER SITES--CONTINUED

ALUM-						MANGA-		
INUM,			IRON,		MANGA-	NESE,		
SUS-	ALUM-	IRON,	SUS-		NESE,	SUS-	MANGA-	_
PENDED		TOTAL	DENDED	TROM				
	INUM,	TOTAL	PENDED	IRON,	TOTAL.	PENDED	NESE,	
RECOV-	DIS-	RECOV-	RECOV-	DIS-	RECOV-	RECOV-	DIS-	
ERABLE	SOLVED	ERABLE				ERABLE		2
								,
(he/r	(he/r			(µG/L		(µG/L	(µG/L	
AS AL)	AS AL)	AS FE)	AS FE)	AS FE)	AS MN)	AS MN)	AS MN	)
392342082072	000 Q-2	SUNDAY C	AT CHAUN	CEY OH (I	AT 39 23	42N LONG	082 07	20W)
0	3,500	54,000	0	54,000	4,400	400	4,000	
385826082201	800 R-1	RACCOON C	AT VINT	ON OH (LA	T 38 58 2	6N LONG 0	82 20 1	8w) 
0	200	480	360	120	1,200	100	1,100	
90	20	790	720	70	1,300	0	1,300	
30	40	870	500	370	360	0	370	
390941082212								v)
20				950			400	
03201988	S-1	L RACCOON	C NR VI	NTON OH (	LAT 38 57	11N LONG	082 21	56W)
700	300	700	550	150	1,900	0	2,000	
270		1,000	850	150	2,000	0	2,000	
0	4,100	430	80	350	4,100	0	4,300	
03160105	S-2	CAMPAIGN				53 51N L	ONG 082	11 31W
80	20	610	480	130	7,100	0	7,300	
383005082280								
	<100		710	150			1,400	
1,340	60	3,500	3,400	130	530	130	400	
120	20	1,300	1,200	100	450	0	450	
382715082242	400 T-2 I	NDIAN GUYA	N C NR B	RADRICK O	H (LAT 38	27 15N L	NG 082	24 24W
	<100	360	220	140	550	100	450	
290	20	890	840	50	760	10	750	
120	40	790	570	220	. 560	0	580	